Division of Geological & Geophysical Surveys

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YUKON RIVER BRIDGE LANDSLIDE: PRELIMINARY GEOLOGIC AND GEOTECHNICAL EVALUATION

by

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Yukon River bridge landslide viewed from the north bank of the Yukon River. Photo by Rich D. Koehler

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YUKON RIVER BRIDGE LANDSLIDE: PRELIMINARY GEOLOGIC AND GEOTECHNICAL EVALUATION

by Rich D. Koehler¹, Richard D. Reger², Karri R. Sicard³, and Eleanor R. Spangler³

Introduction

This report presents the findings of a geologic and geotechnical evaluation of a landslide at the Yukon River bridge, conducted by the State of Alaska, Division of Geological & Geophysical Surveys (DGGS) through a reimbursable services agreement with the State of Alaska Department of Transportation & Public Facilities (ADOT&PF). The E.L. Patton Bridge spans the Yukon River at the Dalton Highway crossing approximately 90 miles northwest of Fairbanks, Alaska (fig. 1). The bridge is a vital link on the transportation corridor between Prudhoe Bay and Fairbanks and supports the trans-Alaska oil pipeline.

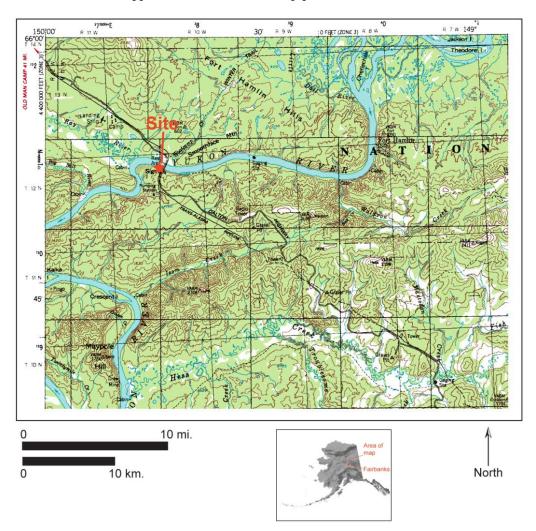


Figure 1. Topographic map of the Yukon River bridge vicinity. Location of landslide shown by red star.

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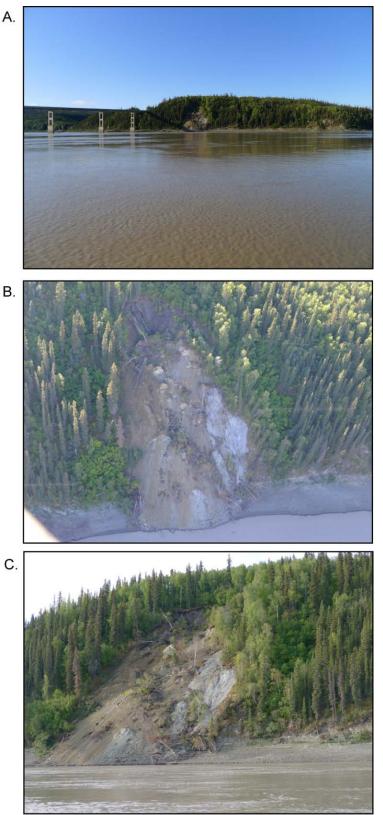


Figure 2. Photographs of the Yukon River landslide. A. View from the north bank of the river. B. Aerial view of slide from helicopter. C. Closer view of slide from the river.

The Yukon River bridge landslide occurred in fall 2012 between approximately 375 and 575 feet west of the bridge (fig. 2). Although there was no damage to the bridge foundation, the landslide's close proximity to the bridge and concerns over additional failures prompted multiple evaluations, including landslide documentation, drainage assessments, and geotechnical studies, among others. The main purpose of this study is to provide baseline geotechnical and geomorphic observations to supplement concurrent slope-stability analyses being conducted by ADOT&PF, Alyeska Pipeline Service Company, and their consultants.

The following tasks are included in the scope of this evaluation:

• Examination of published geologic reports, LiDAR imagery, and geologic maps;

• Aerial reconnaissance of the bridge vicinity by helicopter on August 25, 2013;

• Ground reconnaissance along the base of the bluff below the bridge and around the landslide between August 26 and 29, 2013;

• Preliminary evaluation of rock strength and discontinuity data;

• Preliminary evaluation of principal stress orientations;

• Preliminary evaluation of slopes and landslide geomorphology;

Preparation of this report.

Yukon River bridge and landslide

The E.L. Patton Bridge was originally built in 1974–1975 to facilitate construction of the Trans-Alaska Pipeline System (TAPS). Use and maintenance of the bridge and Dalton Highway were officially turned over to the State of Alaska in 1978. The bridge is a girder bridge design measuring 2,295 feet long and 30 feet wide with a timber deck that slopes upward about 200 feet from the north to the south bank of the river. It is supported by five main in-stream piers, and abutments on each bank. It remains the only bridge crossing of the Yukon River in Alaska. The Dalton Highway and the Trans-Alaska Pipeline System rely on the bridge to connect Fairbanks to the Prudhoe Bay oilfields, ensuring an uninterrupted corridor for the transportation of supplies and personnel. Any compromise to the bridge's integrity could potentially have immediate and severe consequences to the State's economy and environment.

Sometime during late fall 2012, a landslide occurred on the steep bluff above the south shore of the Yukon River between 375 and 575 feet downriver from the E.L. Patton Bridge (fig. 2). The slide originally was reported by Alyeska security on November 30, 2012. Subsequent field inspections conducted by Alyeska and ADOT&PF on December 7, 2012, and January 18, 2013, determined that the slide may have been controlled by highly fractured and altered weak zones in the rock mass (Frank Wuttig, oral commun.). These weak zones were inferred to be related to a potential fault zone encountered during construction of bridge pier 4. Additionally, it was noted that the slide occurred in perennially frozen ground. Given the 2012 slope failure and the occurrence of potentially adverse ground conditions, it was recommended that additional slope monitoring and geotechnical investigations be undertaken. This report contributes to those ongoing investigations.

Regional geology and geomorphology of the Yukon River bridge vicinity

The Yukon River bridge landslide is located along the north-facing bluff of the Yukon River canyon directly west of the E.L. Patton Bridge, approximately 20 miles downstream from the southwestern margin of Yukon Flats (fig. 1). The Yukon River essentially marks the boundary between the Kokrine–Hodzana Highlands, the Rampart Trough, and the Yukon–Tanana Upland physiographic provinces (Wahrhaftig, 1965).

The regional geology of the vicinity of the site is shown in figure 3. Bedrock lithologies near the Yukon River bridge are primarily mapped as Mississippian–Triassic Rampart Group rocks, but minor Tertiary volcanic and sedimentary rocks are present to the southeast (Weber and others, 1992). The Rampart Group consists of a complex assemblage of tectonically deformed mafic igneous and sedimentary rocks. The mafic igneous rocks are intrusive and extrusive and range from aphanitic greenstone to coarse-grained diorite and gabbro, with very rare occurrences of ultramafic rock (Weber and others, 1992). These rocks locally include pillows, amygdules, and interbedded sedimentary rocks. Sedimentary rocks of the Rampart Group, defined by Weber and others (1992), consist of argillite, chert, graywacke, shale, and limestone. Shale units are typically dark gray to black, graphitic, calcareous, and fissile. Argillite and chert units are medium–light gray to greenish gray, locally banded, conglomeratic or brecciated. Graywacke metasandstones are greenish gray, fine to medium grained, and locally micaceous, feldspathic, quartzose, or calcareous. Minor limestone units are bluish gray, fine grained, and laminated. The Rampart Group may correspond to the Tozitna stratigraphic belt in the Livengood area to the south (Dover, 1994; Mertie, 1937; Jones and others, 1984).

In the upland south of the river, tributaries of the Yukon River are incised $\leq 1,700$ feet below ridge crests in rocks of the Rampart Group. Ridge crests are blanketed by 5–50 feet of frozen, locally ice-rich upland loess, and intervening lowlands and stream valleys contain thick, frozen, ice-rich loess and retransported silt (Kreig and Reger, 1982; Shur and others, 2010). These deposits are roughly equivalent to Quaternary loess and colluvium (Qlc) and Quaternary silt (Qsu) mapped by Weber and others (1992) (fig. 3). Through its canyon, the Yukon River winds past 800- to 1,000-foot-high bedrock ridges that were steeply truncated by colossal middle to late Pleistocene outburst floods (Froese and others, 2003; Thorson and Dixon, 1983). At least 160 feet of frozen, fine-grained fluvial and eolian sediments form a surface 200–300 feet above the modern river (Clement, 1999). That surface is modified by thermokarst processes, including formation and expansion of thermokarst pits and gullies, collapse of deep yedoma depressions, and local development of thaw-induced, retrogressive landslides in the river-cut silt bluffs. A pair of late Holocene fluvial terraces with sand and silt cover discontinuously flanks the modern river channel.

The Ray River lowland contains a thick, perennially frozen fill of ice-rich, fine-grained sediments that accumulated during massive late-Pleistocene outburst floods in a major slackwater basin just upriver of a hydraulic dam that formerly existed where the flooding Yukon River abruptly turned south and entered a narrow canyon (Thorson, 1989). Northeast of the Ray River lowland the Fort Hamlin Hills, which are composed of Rampart Group that is weathered to depths \geq 5 feet, rise to elevations of 1,500 to 3,150 feet. Up to one foot of loess discontinuously covers bedrock and colluvium on upper slopes, and frozen, locally ice-rich loess and retransported silt reach thicknesses of \geq 100 feet on middle and lower slopes. Perennially frozen alluvial–colluvial fans coalesce at the mouths of stream drainages cut into the southern flank of the Fort Hamlin Hills, forming an ice-rich apron crossed by the TAPS and Dalton Highway routes.

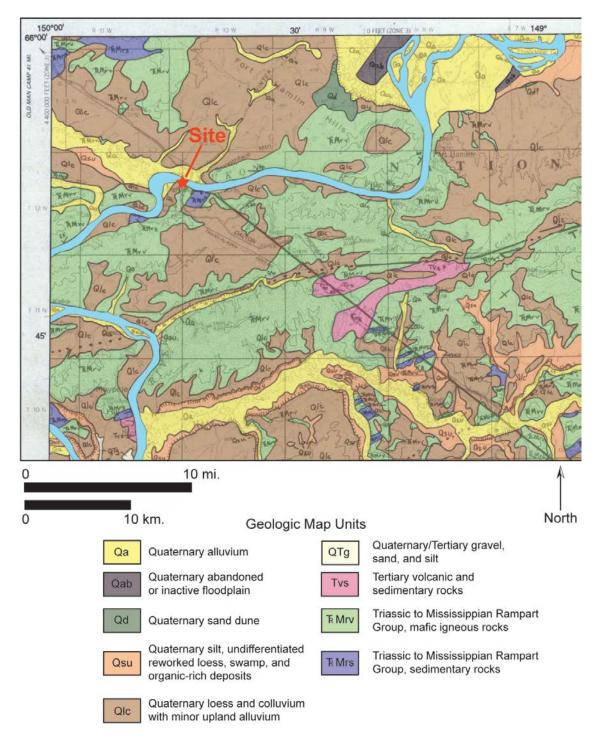


Figure 3. Geologic map of the site vicinity from Weber and others (1992). Geologic units are colored to emphasize unit boundaries. Location of the Yukon River bridge landslide is shown by red star.

Yukon River bridge landslide field review

Methodology

The Yukon River bridge landslide field review was performed August 25–29, 2013, by Rich Koehler, Karri Sicard, and Eleanor Spangler of DGGS and Richard "Dick" Reger of Reger's Geologic Consulting. A helicopter reconnaissance organized by Frank Wuttig of Alyeska Pipeline Services Company provided an overview of the landslide and general observations of the slope conditions upstream and downstream of the bridge. Robert Joseph of Stevens Village provided boat transportation to access the lower bluff and landslide area and to observe slopes along the south bluff upstream and downstream of the bridge.

Bedrock outcrops along the base of the Yukon River bluff were inspected for the presence of faults, fractures, shears or weathering zones, and/or other features that could provide information from which to infer subsurface conditions of bedrock farther upslope. Quantitative descriptions of rock mass and discontinuity data are based on the methodology of ISRM (1978). Rock mass and discontinuity survey data sheets from the Alaska Field Rock Classification and Structural Mapping Guide (ADOT&PF, 2003) were used to record the field observations. Key rock characteristics that were documented include lithology, color, grain size, field compressive strength, rock mass fabric, block size, and state of weathering. Discontinuity surveys at each outcrop station include information on type of discontinuity, dip, dip direction, persistence, termination, aperture, nature of filling, strength of filling, surface roughness, surface shape, waviness, joint roughness coefficient, and presence of water. Detailed definitions for each of these parameters are provided in ADOT&PF (2003). Discontinuity orientations were measured with a Brunton compass, using standard geologic field techniques.

The slope and landslide evaluation made use of both office and field techniques. Pre- and post-landslide LiDAR data collected by Hubbard and others (2011) and Alyeska Pipeline Service Company, respectively, were used to generate slope profiles and evaluate possible past slope failures. These two data sets were also used to compare the pre- and post-landslide ground surface. The combined LiDAR and field observations were used to quantitatively evaluate possible failure mechanisms for the 2012 landslide.

Results of rock mass and discontinuity evaluation

We evaluated rock strength and discontinuity characteristics along the base of the bluff and riverbank at 13 bedrock field stations distributed between approximately 850 feet east of the Yukon River bridge and approximately 1,150 feet west of the bridge (fig. 4). An additional zone of intensely sheared graywacke between stations 10 and 11 was mostly covered by beach gravels and boulders, and thus not described in the field as a separate station. Observations from this zone are provided in the following discussion. Photographs documenting bedrock features at each outcrop station along the base of the bluff are included in Appendix A. An additional observation station was located at the crest of the landslide. The location of each station and approximate bedrock contacts are shown on a strip map of the transect (fig. 4). GPS coordinates for each station are listed in Table 1. Detailed observations and inventory of rock discontinuity measurements are recorded in the field evaluation sheets in Appendix B.

Bedrock lithologies along the base of the Yukon River bluff consist of diorite, gabbro, graywacke metasandstones, greenstone, metabasalt, and chert (fig. 4). Table 2 shows rock-mass information for each bedrock field station. In general, the rock mass is characterized as strong to very strong rock that is fresh to slightly weathered with local zones of moderate weathering. The freshness of the rock is likely a reflection of ice scour and limited exposure to surface conditions (i.e., <15,000 years.). Limited inspection of the roadcut at the south abutment of the bridge indicates that the rock may grade upslope to a more weathered condition (moderately to highly weathered?); however, this outcrop was not studied in detail. The rock character along the base of the bluff is generally blocky (small to medium in scale) with discontinuous zones of very small shattered blocks and closely fractured to crushed zones. The graywacke metasandstone is the only rock type with a discernible fabric, including slightly metamorphosed and deformed beds. However, the deformed bedding is irregularly folded and does not have a consistent orientation.

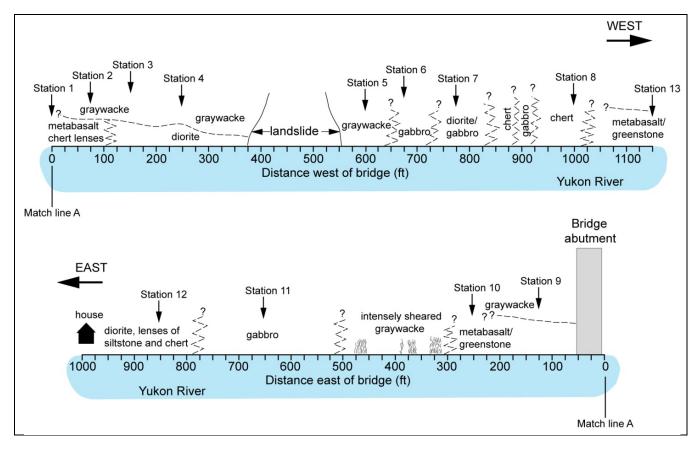


Figure 4. Location of field stations and inferred bedrock contacts along the Yukon River bluff transect. Area west of the bridge is shown at top; area east of the bridge is at bottom.

Station	GPS			
number	point	Latitude	Longitude	
1	5	65.87386703	-149.724802	
2	6	65.87396703	-149.725185	
3	7	65.87400097	-149.725785	
4 (a and b)	8	65.874034	-149.725752	
5 (a and b)	9	65.87451696	-149.728152	
6	10	65.87496699	-149.728452	
7	11	65.87491703	-149.729318	
8	12	65.87515097	-149.730452	
9	13	65.87381699	-149.724118	
10	14	65.87370099	-149.723002	
11	15	65.87271703	-149.721452	
12	16	65.87230104	-149.720385	
13	18	65.87573804	-149.732431	

Table 1. Locations of data stations. GPS data collected in WGS 84.

Station	Lithology	Color	Grain size	Compressive strength	State of weathering	Fabric	Block size	Number of discontinuity sets
1	1 Greenstone/ Dark greenish metabasalt gray to grayish green		Very fine	R5	Fresh	Blocky	Very small	3–4
2	Metabasalt	Dark greenish gray	Fine to very fine	R5	Fresh	Blocky	Medium	>5
3	Diorite	Reddish black to green	Fine	R5	Slight	Blocky	Medium	5
4a	Diorite	Dark reddish gray	Fine	R4	Moderate	Blocky	Small	5
4b	Graywacke metasandstone	Light bluish gray	Very fine	R5	Slight	Blocky to shattered	Small	4
5 (a and b)	Graywacke metasandstone	Dark greenish gray to light green	Fine to very fine	R4-R5	Fresh to slight	Blocky, columnar, shattered	Medium	3
6	Gabbro	Light brownish gray	Fine	R5	Fresh to slight	Blocky	Medium	>4
7	Diorite to gabbro	Dark greenish gray	Fine	R5	Slight	Blocky	Medium	5
8	Chert	Brownish red to greenish blue	Very fine	R4	Slight	Shattered	Very small	3
9	Metabasalt	Dark reddish to yellowish gray	Very fine	R4	Slight	Blocky	Medium	2
10	Basalt/ greenstone	Dark greenish gray	Very fine	R5	Slight	Blocky	Medium	4
Area between 10 and 11	Graywacke metasandstone	Greenish gray to olive green	Fine	R1-R4	Slight	Shattered	Very small	3
11	Gabbro	Dark grayish black	Fine	R5	Moderate	Blocky	Medium	3
12	Diorite	Dark bluish gray	Fine	R5	Slight	Blocky	Small	3
13	Metabasalt/ greenstone	Dark greenish black	Fine	R5	Slight	Blocky	Small	4

Table 2. Rock mass data for individual measurement stations.

Discontinuity data evaluation

Jointing is pervasive throughout the rock mass for all bedrock types and, together with local faults and shear zones, divides the rock into discrete blocks and slabs. Joint dips are predominantly high angle (Appendix B) and may have formed during uplift or unroofing of the Yukon–Tanana Upland. Joint spacing is typically close to moderate in all rock types except chert, which exhibits extremely close to very close spacing possibly related to bedding-parallel shearing, or due to the very fine grain size. Joint surfaces are typically tight and smooth to slightly rough and planar with scattered slickensided shear surfaces. Scattered black oxidation staining and thin fillings of calcite and quartz are present on joint surfaces, however the majority of joints are clean. Aperture widths are commonly tight to very tight, with scattered joints partly open to open and rare very wide openings. Evidence of water flow was not observed on any discontinuity surface.

A wide area of sheared chloritic graywacke metasandstone with zones of more intense shearing and alteration is present between approximately 300 and 475 feet east of the Yukon River bridge (fig. 4). The graywacke between shear zones is strong rock (R4). The shear zones are characterized by very weak (R1), shattered and brecciated rock in a grayish green clay matrix with calcite veins. Shear zones are variably oriented with clusters striking 60° – 90°

azimuth and 145° – 170° azimuth with vertical dips. The zones are >18 inches wide, closely spaced, and irregularly curved with calcite present as skins on shattered clasts. Individual shear planes in the zones are tight to open and have orientations similar to the major shear zones.

The 13 bedrock field stations are characterized by an array of joint and fault orientations. Rose diagrams and equal-area stereonet projections of the discontinuity data were plotted using the program Stereonet 8 (Allmendinger and others, 2012). Figures 5 and 6 show rose diagrams of joint orientations for each rock type and field station, respectively. The measurements were reflected across the rose diagrams for visual ease, so that both directions of each plane are visually represented. The bin size is 10 degrees, and the discontinuities are not weighted by length, persistence, spacing, or other factors.

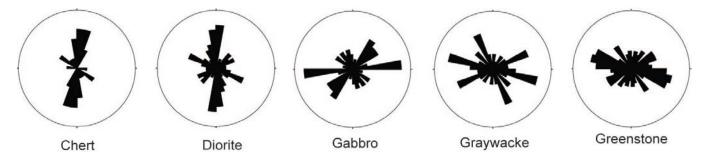


Figure 5. Rose diagrams of rock discontinuity data by rock type. Measurements are reflected across the rose diagrams so that both directions of each plane are represented.

Discontinuity orientations for the different rock types (fig. 5) show clear distinctions. Chert and dioritic rocks have similar fracture patterns with dominant N–S joint sets, and subdominant WNW–ESE and NE–SW fracture sets. These rock types are both fine grained, and the fractures may have broken the rock without being influenced by crystal size or pre-existing weaknesses. Fracture patterns in the gabbroic rocks are characterized by NE–SW and E–W fracture sets. Graywacke metasandstones are characterized by three dominant sets, NW–SE, NE–SW, and WNW–ENE. Pre-existing weaknesses related to bedding in the graywacke metasandstones may have been exploited by the fractures. Greenstones (metabasalts) have fracture patterns that are similar to the dominant WNW–ESE set in the graywacke metasandstones, but also contain NW–SE and NE–SW fracture sets. Shear zones and faults in the greenstones may have influenced the discontinuity orientations. The maxima of the discontinuities are WNW–ESE (~285°–95° azimuth), N–S (~355°–175° azimuth), and NE–SW (~30°–210° azimuth).

Additional patterns are evident in the rose diagrams for individual field stations (fig. 6), however because each station has only about 20 measurements, the patterns may reflect uneven sampling. The rose diagram for the combined dataset of all stations includes 610 discontinuity measurements (fig. 6). This plot generally shows the same major discontinuity sets as seen in the different bedrock type plots but is more scattered, illustrating the diverse complex fracture pattern that characterizes the site.

The strike and dip of the discontinuity surfaces and poles to discontinuity planes were also plotted on equalarea stereonet projections (fig. 7). For the entire data set, no distinct patterns are evident in the composite stereonet projections of the strike and dip measurements (fig. 7A). As with the rose diagrams, grouping stereonets by station and rock type may reveal more robust patterns, however, this is beyond the scope of our study. The plot of the poles to the discontinuity planes shows several distinct clusters in the main populations (fig. 7B). The lines around the clusters are 1 percent area contours. The main plane orientations dip nearly vertical (scattered toward the edges of the net) and are similar to the rose diagrams; the dominant WNW–ENE planes are the cluster of poles plotting slightly E of N, and the dominant N–S planes are the cluster of poles near the E and W edges of the net. The NE–SW planes do not show up as a large cluster of poles, but there are small clusters dipping less vertically (~75°) in the NW and SE quadrants. The last cluster of poles, approximately 45° E of N near the plot edge, is the near-vertical NW–SE set of discontinuity planes.

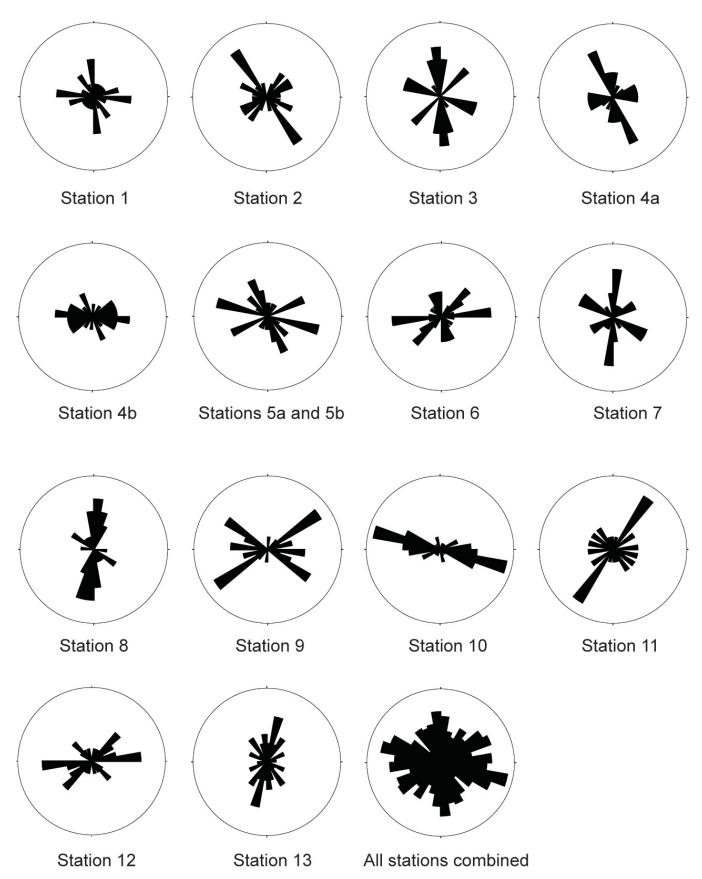


Figure 6. Rose diagrams of rock discontinuity data from each measurement station, and composite plot of the combined dataset (lower right). Measurements are reflected across the rose diagrams so that both directions of each plane are represented.

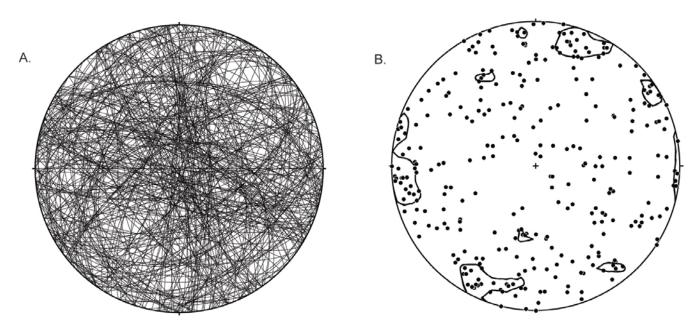


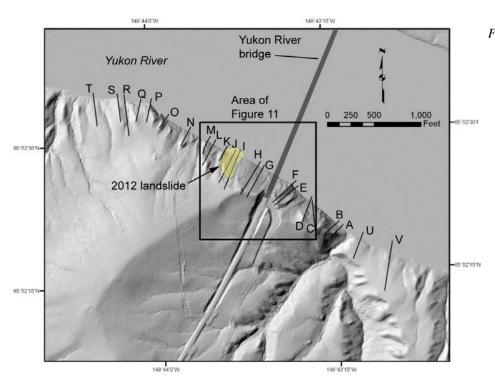
Figure 7. Stereonet projections of rock discontinuity data. A. Equal area stereonet projections for all discontinuity data.B. Poles to the discontinuity planes for all discontinuity data. Lines around clusters are 1 percent area contours.

There are additional ways to analyze the discontinuity data, but these studies should be driven by the goals of future slope stability evaluations and are beyond the scope of this evaluation. For example, discontinuities could be weighted based on important factors such as their length, persistence, spacing, oxide staining, fillings, and other characteristics. The discontinuities could be evaluated spatially, which could reveal patterns related to proximity to shear zones or large faults. Different blocks may exist in areas with different stresses acting on them currently or in the past. Additionally, the fractures proximal to the landslide could be compared to more distal fractures to evaluate whether there are more-dominant fracture sets that could have controlled the bedrock failure in the area of the landslide. A preliminary test of this hypothesis did not yield patterns.

Fault evaluation

Evaluation of fault orientations can provide insight into the state of stress at the time of faulting. Large datasets are needed to determine the current and past stress fields, since stress usually changes through time, and records of these different states of stress are recorded in the rocks. Cross-cutting relations and mineral growth on fault surfaces are extremely useful for evaluating these stresses. Also, conjugate fault sets are much more instructive than independent fault planes in evaluating and determining the state of stress. The pattern of slip on a simple conjugate array of mesoscopic faults is directly related to the state of stress at the time of faulting.

We measured 37 mesoscopic faults in the area around the Yukon River bridge, 33 of which had shear fractures with slicken surfaces. The lack of openings, water flow, gouge, and damage zones on the majority of these faults indicates that they are likely old, and no evidence of recent movement was observed. Some of the shear surfaces showed two different sets of slickenlines with different sets of motion, indicating that the movement had changed. The slip planes of some of these features appeared to be along curviplanar quartz and sometimes epidote veins (plus or minus chlorite), and were sealed or very tight. This suggests that the tensile openings were exploited as slip surfaces. Conversely, the veins could have taken advantage of the fault openings and slickenlines may have been created by continued slip. Either way, the sealed nature of the fractures suggests that this tensile state has passed. Further fault analyses could elucidate the sequence of past structural deformation in the area, however it would not help determine the current state of stress. Thus we did not analyze the faults separately.



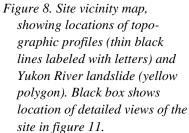


Figure 9. Topographic profiles A–H and K–V across the southern bluff of the Yukon River in the vicinity of the Yukon River bridge. Figure 8 shows the location of each profile. Profiles I and J across the Yukon River bridge landslide are shown in figure 12.

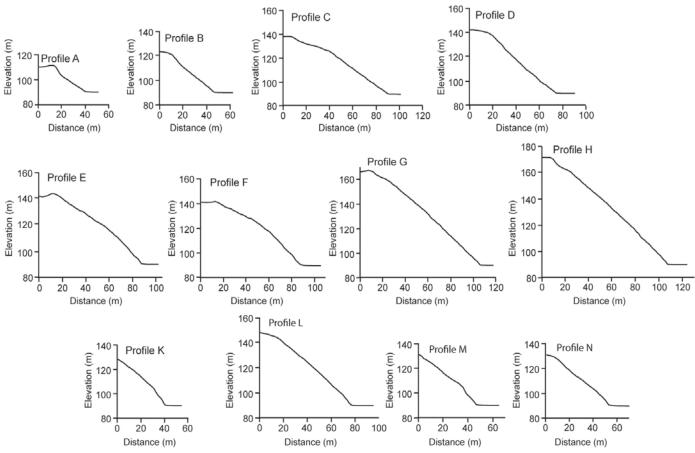
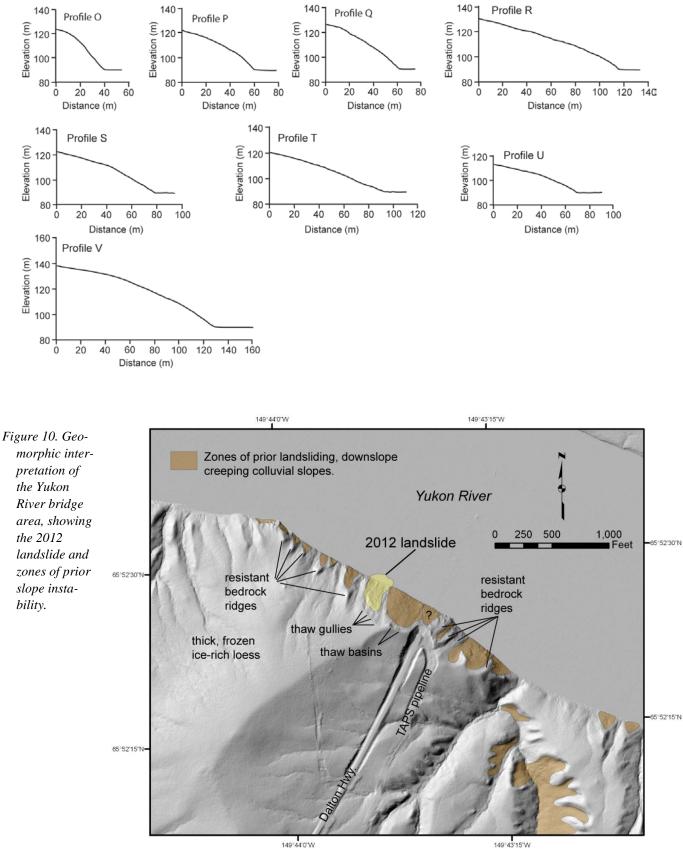


Figure 9, continued



149°44'0'W

Results of slope and landslide geomorphology evaluation

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General slope conditions and the location of individual slope profiles are shown in figure 8; profiles are in figure 9. A LiDAR interpretation of the location of prior landsliding and/or other downslope colluvial transport processes is shown on figure 10. For the areas upstream and downstream of the bridge, LiDAR imagery of the bluff shows several vegetated possible landslide slopes separated by steep, narrow bedrock ridges that are relatively stable (fig. 10). Surfaces above the crest of the bluff slopes are mantled by thick, frozen ice-rich loess and characterized by gentle slopes and common thaw gullies and basins. Geomorphic features indicative of active faulting are absent on the surface. In particular, there is no indication of recent along-strike activity of the fault identified in the foundation of bridge pier 4. In the field, helicopter surveys and boat reconnaissance were conducted along the bluff to look for evidence of prior landsliding for approximately 2,000 feet upstream and downstream of the bridge; however, no scarps or obvious debris hollows were observed.

Profile						
designation Evaluation scale		Comments				
A	1:398	Slope planar @ 35.5°. Slow creep/gelifluction likely.				
В	1:521	Slope planar @ 38.5°. Slow creep/gelifluction likely.				
С	1:565	Slope planar @ 36.5°. Slow creep/gelifluction likely.				
D	1:610	Slope planar @ 38°. Slow creep/gelifluction likely.				
E	1:530	Steep gully between bedrock ridges contains slope debris. Possible small failure below 324 ft elevation due to river undercutting.				
F	1:549	Steep nose of bedrock ridge. Lower section appears oversteepened by river erosion. Possible slope failure below ~40 ft elevation.				
G	1:719	Possible slope failure below ~539 ft elevation.				
н	1:698	Possible slope failure below ~543 ft elevation.				
I	1:693	2012 landslide below ~507 ft elevation.				
J	1:273	2012 landslide below ~477 ft elevation.				
К	1:278	Steep nose of bedrock ridge. Possible slope failure at ~325 ft elevation due to undercutting by river.				
L	1:522	Uniform slope of 40°. Possible slope failure below ~332 ft elevation due to undercutting by river.				
М	1:347	Steep (40°) nose of bedrock ridge. Possible slope failures below ~339 and ~317 ft elevation.				
Ν	1:348	Steep (40°) nose of bedrock ridge. Steepened to 43° below ~325 ft elevation.				
0	1:348	Steep (45.5°) nose of bedrock ridge. Top of possible landslide at ~334 ft elevation.				
Р	1:523	Margin of possible dissected, old landslide below ~372 ft elevation in hollow west of prominent bedrock point.				
Q	1:521	Axis of possible dissected, old landslide with top at ~385 ft elevation and surface slope of 28°. Slope steepens to 44° below ~338 ft elevation due to river erosion. Top of possible younger landslide at ~321 ft elevation.				
R	1:521	Gully fill between bedrock ridges slopes 17°. Slope steepens to 31.5° at ~337 ft elevation in response to river erosion.				
S	1:521	Loess-covered bedrock ridge. Nose slopes consistent 28° below ~361 ft elevation.				
Т	1:539	Loess-covered bedrock ridge slopes consistently at 20° below ~358 ft elevation.				
U		Not evaluated.				
V		Not evaluated.				

Table 3. Characteristics of LiDAR-derived slope profiles evaluated for evidence of slope instability.

Profiles derived from LiDAR data demonstrate that vegetated possible landslide surfaces slope between 35° and 42° and the steep, lower crests of resistant bedrock ridges slope approximately 40° to 68°. Table 3 shows the scale at which slopes were evaluated and the slope interpretation for each profile. Between rock exposures on unfailed slopes, 2–3 feet of loess overlie 4–6 feet of pebbly, silty colluvium derived from downslope transport of loess and weathered bedrock. Slopes adjacent to the 2012 landslide display surface evidence of downslope movement,

including numerous bent trees and shrubs, solifluction lobes, and scattered 10- to 20-foot-long, gaping transverse cracks. Profiles A through H and K through Q show oversteepened bases that may be a result of river erosion or past massive floods and/or ice scour (fig. 9). These profiles are located in the zone of possible previous slope movements shown on figure 10. Although some of these slopes may have catastrophically failed in mass wasting events, they may also reflect long-term downslope creep and gelifluction of weathered (residual) bedrock and loess. The absence of large colluvial hollows and debris blocks, as well as the relatively planar morphology of the slope suggests that the latter is the more dominant process of slope modification. Profiles R through V show more subdued profiles and are generally west and east of the zone of possible slope movements. In these areas, the slopes are not faceted and appear relatively stable.

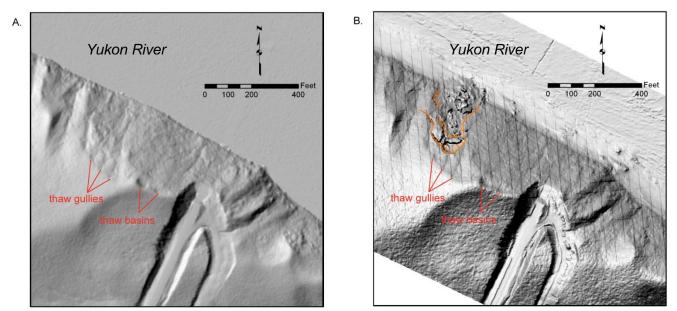


Figure 11. Pre- and post-landslide hillshade images derived from 2010 and 2013 LiDAR data. A. Pre-landslide hillshade. B. Post-landslide hillshade; edges of landslide scarps shown by orange lines.

A comparison of the pre- and post-landslide surface and profiles across the landslide are shown in figures 11 and 12, respectively. Photographs of key features on the landslide are shown in Figure 13. In the pre-landslide hillshade image, hummocky topography along the slope indicates the possible presence of previous slope instability. Thaw gullies at the crest of the slope may have contributed to previous slope movements. The 2012 landslide occurred along the downstream end of the thaw gullies. Based on the distribution of headscarps (orange lines on fig. 11B), the 2012 landslide initiated on the east and propagated to the west. Blocks of debris in the mid-slope region and a large debris cone at the base of the slope are clearly visible on the LiDAR. Based on the LiDAR data, the main slide mass, including the area of the debris cone, is approximately 90–100 feet wide and approximately 220 feet long, however the bulk of the debris was sourced from an approximately 90×100 -foot translational slide block in frozen loess at the top of the slope (fig. 13B). In addition to the frozen loess, a large slab of highly weathered gabbro also failed. Also present are two additional blocks of rock and colluvium that did not catastrophically fail down the slope. An approximately50x50-foot block on the western side of the headscarp separated from the slope across a large fracture. This block is underlain by fairly competent, fractured graywacke metasandstone and was likely made unstable after the main slide eroded the downslope side of the block, removing lateral support and causing slip to occur along fracture planes. Midway along the western margin of the slide a second block, measuring approximately 50 × 100 feet, separated approximately three feet from the adjacent inplace slope (fig. 13D). This block may be susceptible to future sliding.

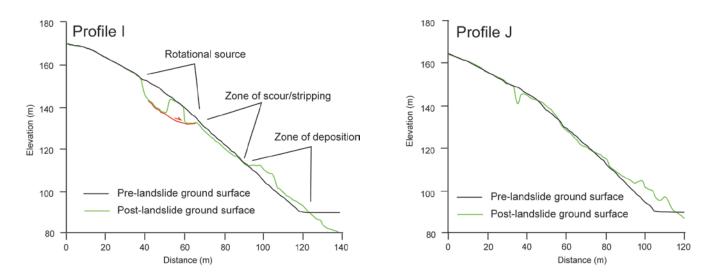


Figure 12. Topographic profiles I and J across the Yukon River bridge landslide showing the pre- and post-landslide ground surface. Pre-landslide ground surface derived from LiDAR data collected by DGGS (Hubbard and others, 2011). Postlandslide ground surface derived from LiDAR data collected by Alyeska Pipeline Service Company in April 2013. Inferred slide plane indicated in orange on Profile I.

Topographic profiles I and J were generated from the pre- and post-landslide LiDAR data across the eastern and western sides of the body of the 2012 landslide. They demonstrate that the crown of the slide is approximately 236 feet above low-water level (fig. 12). The profiles show that the landslide occurred on a slope of approximately 40° and that the post-landslide slope is around 35–37°, which was confirmed by direct measurements in the field. Profile I, down the main body of the landslide, shows three distinct zones including an upper block failure, a mid-slope zone of scour and stripping, and a lower zone of slide debris deposition (fig. 12A). Flood deposits on the lower part of the debris cone indicate that break-up floods in spring 2013 did not remove an appreciable amount of slide debris. Profile J, measured on the slope directly west of the main slide body, shows a large tension fracture and only subtle translation of the slope mass (fig. 12B).

Three 20- to 30-foot-wide former thaw gullies are present on the slope across the top of the landslide, but their rounded cross profiles indicate that they are filled with loess and retransported loess and are no longer active. Along the upper marginal scarps and headwall of the failure, the loess displays closely-spaced planar structures that generally parallel the ground surface. In Interior Alaska, such planes develop in loess during the annual freezing of the ground by the formation of very thin, clear, parallel to subparallel, segregated ice lenses (French, 2007). Parallel shear planes also develop parallel to the slope during slow downslope creep of the loess, and downslope creep likely produced the parallel shear planes in the underlying frozen colluvium. A 5-inch-wide wedge of clear ice was visible for approximately 5 feet across the freshly exposed headwall of the landslide, demonstrating that the permafrost table was approximately 2 feet deep at the end of August 2013. The ice wedge formed when precipitation or meltwater entered and froze in an open transverse crack on the perennially frozen slope. No other visible ice was observed in the permafrost there, and no evidence of meltwater from thawing permafrost was observed in or around the landslide.

Field observations indicate that the large landslide blocks of frozen loess, colluvium, and highly weathered bedrock were not rotated during downslope motion but slid on the underlying, more competent bedrock on a shallow, planar, translational slide plane. Northwest-trending transverse open cracks and their ice fillings likely had a significant impact on the locations and sizes of the block failures. Additionally, long-term creep and gelifluction on the slope may have contributed to progressively greater tension that eventually exceeded the tensile strength in the frozen blanket of loess and colluvium. Deeper slide planes may also have developed along

intersecting fracture planes in the highly weathered gabbro, which is altered to seamy clay. After initial failure, the displaced blocks collapsed and cascaded down the steep slope as a debris slide, stripping 2–3 feet of colluvium off the underlying bedrock in the mid-slope region, and depositing a large debris cone over the lower 90 feet or so of the slope. The debris cone is characterized by angular chert, gabbro, and graywacke metasandstone fragments, averaging 4–18 inches in diameter. Several large debris blocks with trees in growth position were deposited on the slope (fig. 13C). The zone of stripping and talus deposition is shown in figure 13A. Although a more detailed site investigation is necessary to more accurately define the slide plane and mechanism of failure, the field observations are consistent with a block glide transitioning downslope to a debris slide (Varnes, 1958).

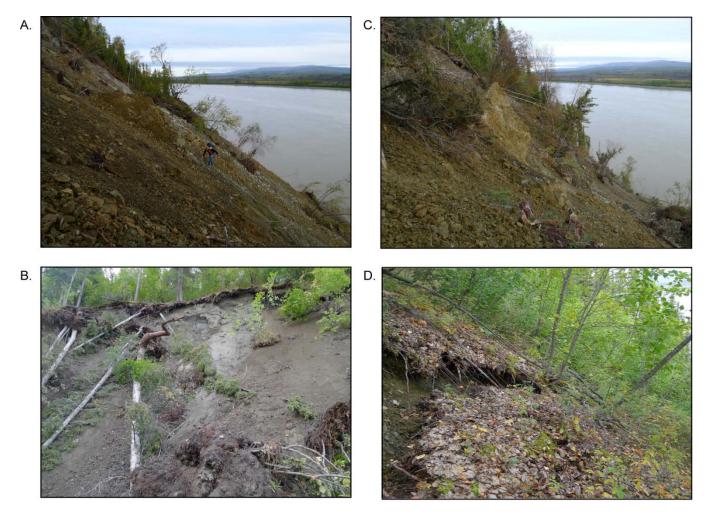


Figure 13. Photographs of the Yukon River landslide. A. Upper part of debris cone. Light gray color in upper left is exposed graywacke bedrock. B. Intact block of weathered gabbro bedrock on the surface. C. Frozen silt exposed in the headscarp. D. Separation scarp extending west of the middle part of the slide.

Conclusions and Recommendations

The 2012 landslide occurred adjacent to the Yukon River bridge, but the bridge foundation did not sustain any damage. The slide initiated in frozen loess and retransported loess, but also involved the underlying highly weathered, weak bedrock (gabbro and graywacke metasandstone). The high density of fractures, joints, shear zones, and minor faults in the rock mass likely contributed to instability, and active faulting was not a factor. Tension in the overlying loess related to transverse cracks and ice fillings and the influence of surface hydrology may also have contributed to the slope failure. The relative roles of these potential causative factors are not completely understood. Based on the surface geomorphology, we conclude that the slide is most consistent with a block glide failure transitioning downslope to a debris slide.

The bedrock and discontinuity survey determined that the bedrock along the base of the bluff is primarily composed of metabasalt (greenstone), diorite, gabbro, graywacke metasandstone, and chert with minor limestone, all of which is fresh and strong. The rocks are closely jointed, with dominant joint sets oriented WNW–ESE, NE–SW, N–S, E–W, NW–SE. These joint patterns likely extend throughout the rock mass, which includes zones that are highly altered and intensely weathered upslope. Our assessment of slopes in the vicinity of the bridge indicates that slopes are >40° and are mantled by fairly thin colluvium. The crest of the slope is buried by thick frozen loess. The interaction of gelifluction and ice and soil processes with the intensely fractured rock provides favorable driving forces for slope failure.

Several arcuate, open fractures are present south of the main headscarp of the 2012 landslide, indicating that additional smaller landslide blocks are poised for future downslope displacements. The slightly separated slope along the western margin of the landslide is prone to future failure, assuming continued loss of lateral support and steepening of the base of the slope by Yukon River erosion. Thaw basins at the top of the slope between the landslide and the Yukon River bridge, as well as the presence of gelifluction and active creep features, indicate the possibility of similar slope failures occurring in that area in the future. Although the potential exists for future slope failures, the lack of any significant landslide scars along the bluff upstream and downstream of the bridge indicates that large-scale failures are a rare occurrence. The 2012 landslide is the first landslide of this type and scale at least since the construction of the Yukon River bridge in 1974–75 and probably for at least 20 years previously.

Our geologic and geomorphic evaluation has determined that the fractured bedrock and steep slopes in the bridge vicinity are prone to slope failure. Given the significant landslide hazard adjacent to the Yukon River bridge and Trans-Alaska Pipeline System, and the potential impact of damage to the bridge to the state's economy, it is recommended that a monitoring and instrumentation program be initiated on the slope adjacent to the bridge to better understand the landslide risk. The instrumentation program should include the installation of piezometers, extensometers, and tiltmeters. Detailed hydrologic studies aimed at characterizing drainage patterns, and a geotechnical study focused on characterizing subsurface conditions and an estimate of "Factor of Safety" are also recommended, particularly in the critical area between the 2012 landslide and the Yukon River bridge.

Closure and Limitations

This report was prepared to convey to the public the general characteristics of the rock mass, characteristics of rock discontinuities, and the geomorphic expression of the 2012 landslide in the vicinity of the Yukon River bridge. The observations and conclusions contained in this report are based on site conditions on the dates of the field evaluations discussed herein and are the opinions of the authors. The information contained in this report should be considered preliminary and should not be used to determine areas of potential future slope instabilities. A significant amount of additional exploration and slope stability analysis are required to fully understand the landslide hazard in the area of the bridge, as well as the potential for reactivation and/or expansion of the 2012 landslide into adjacent slopes.

Acknowledgments

The authors appreciate the coordination among Steve McGroarty, Jeff Currey, Kevin Maxwell, and Garrett Speeter of (ADOT&PF) and Frank Wuttig of Alyeska Pipeline Services Company at pre-project briefing and planning meetings as well as during our field evaluation. Alyeska graciously provided helicopter support for our initial reconnaissance. Robert Joseph of Stevens Village expertly guided our boat-based river surveys. Eric Hatleberg of DNR's State Pipeline Coordinator's Office provided valuable feedback regarding our work during a field review. We thank De Anne Stevens of DGGS for an insightful review of the report.

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Appendix A

Photographs of bedrock at rock discontinuity measurement stations

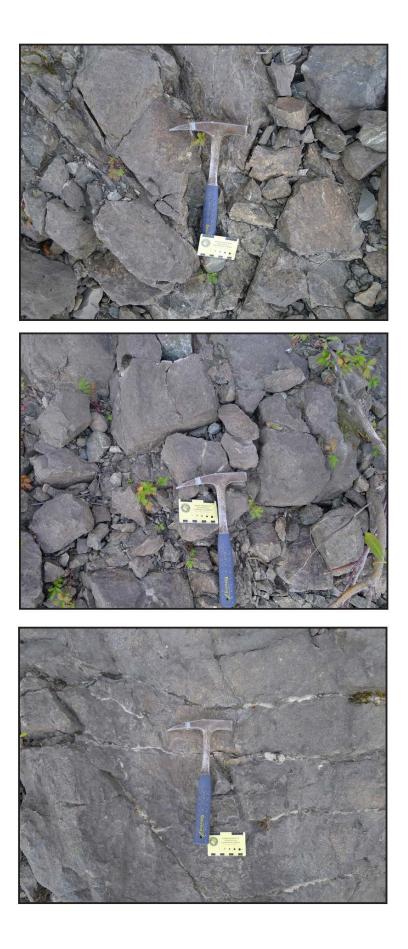
A1. Station 1 joint sets, beneath the Yukon River bridge.





A2. Station 2 joint sets, 75 feet west of the Yukon River bridge.

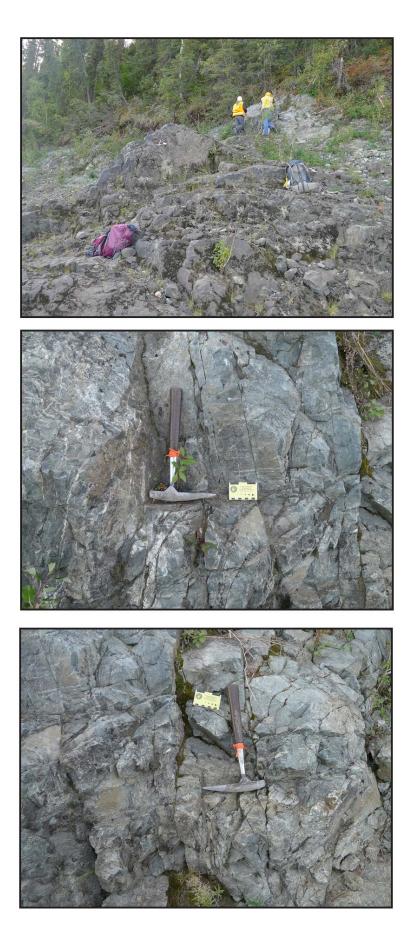
A3. Station 3 joint sets, 150 feet west of the Yukon River bridge.





A4. Station 4a joint sets, 250 feet west of the Yukon River bridge and directly east of landslide.

A5. Station 4b joint sets, 250 feet west of the Yukon River bridge and east of the landslide.





A6. Station 5a joint sets, 600 feet west of the Yukon River bridge and west of the landslide.



A7. Station 5b joint sets, 600 feet west of the Yukon River bridge and west of the landslide.



A8. Station 6 joint sets, 650 feet west of the Yukon River bridge and west of the landslide.

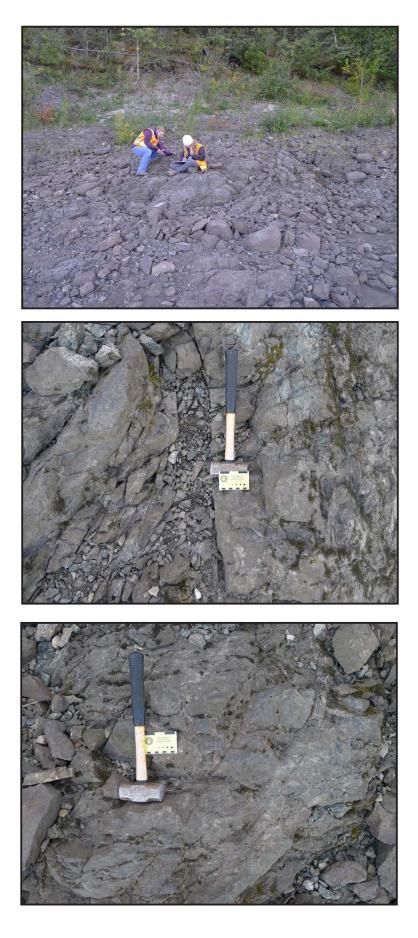


A9. Station 7 joint sets, between 770 and 800 feet west of the Yukon River bridge and west of the landslide.



A10. Station 8 joint sets, 1,000 feet west of the Yukon River bridge and west of the landslide.

A11. Station 9 joint sets, between 100 and 150 feet east of the Yukon River bridge.





A12. Station 10 joint sets, between 200 and 250 feet east of the Yukon River bridge.

A13. Station 11 joint sets, between 625 and 675 feet east of the Yukon River bridge.





A14. Station 12 joint sets, between 850 and 900 feet east of the Yukon River bridge.

A15. Station 13 joint sets, approximately 1,150 feet west of the Yukon River bridge (promiment outcrop that extends into the river).

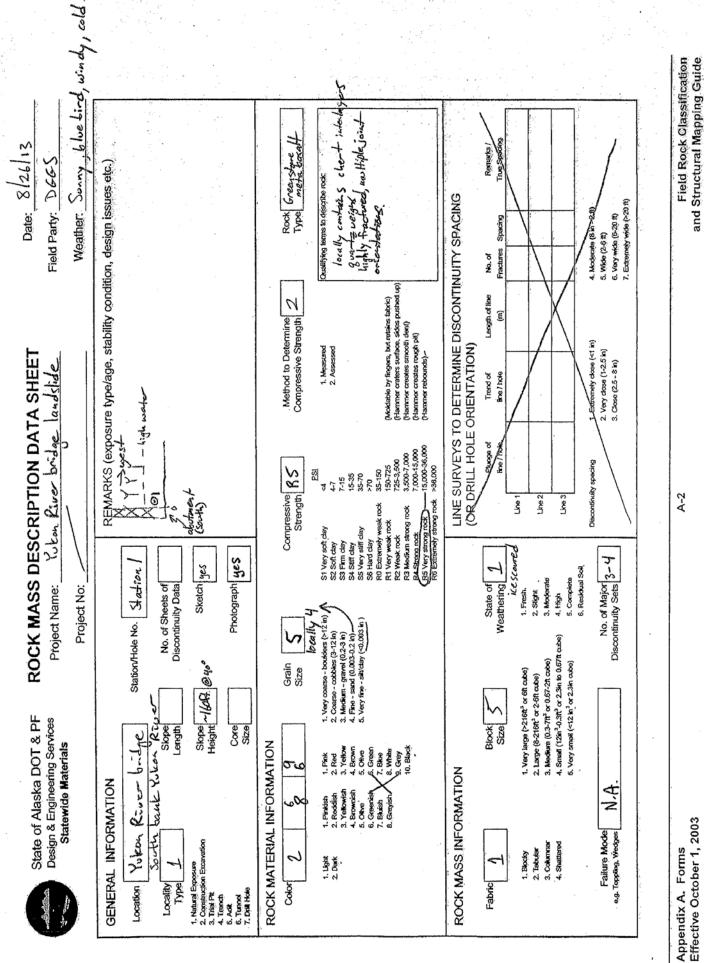




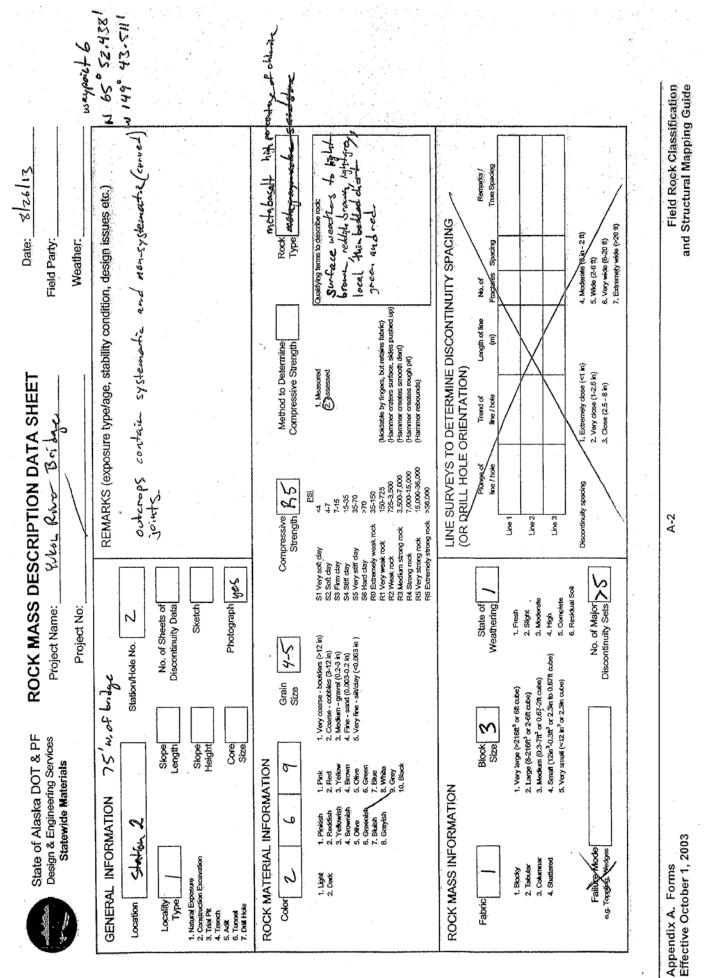
A16. Area between Stations 10 and 11, joint sets in sheared graywacke, approximately 300 to 475 feet east of the Yukon River bridge.

Appendix B

Field evaluation sheets

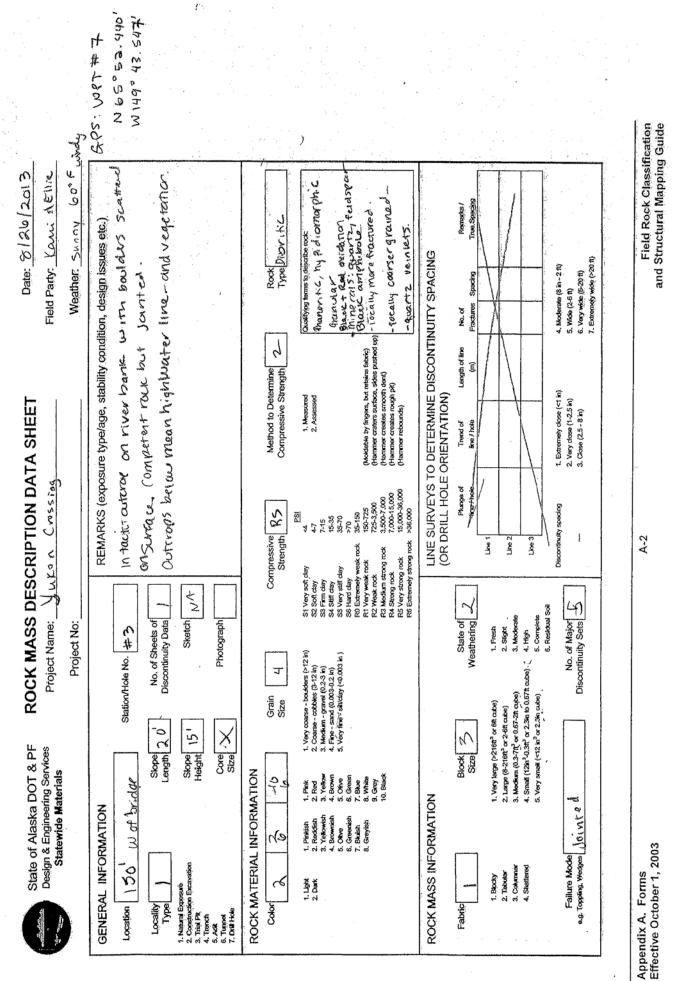


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Yukon River bridge landslide: Preliminary geologic and geotechnical evaluation

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)OT & P og Service erials			N OF DI		Persistence	n¢	d	2	-	d a	20	d	6	2	d	R	2	2	. 1	1	ሪ	ત	ർ			- 5		7. Vecy 8. From	9. Caw						
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Appendix A. Forms Effective October 1, 2003

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Field Rock Classification And Structural Mapping Guide Sheet 10f2

		e de la com				
Date: <u>5/27/2013</u> Field Party: <u>Lawi + EUi e</u> Weather	e type/age, stability condition, design issues etc.) a.S. to the E of the 250' Station and if the Stathon	8	Rock Dorth C. Type Dorth C. Fuendrin Back anghr hald, guartz, Readort C. Kypickingher hald, guartz, Readort C. Verns Preve kent eboart C. Verns Preve kent Verns. Verns.	UITY SPACING No. of Remeter/ Fractures Spacing True Spacing		4. Moderate (8 in - 2 ft) 5. Wide (2.6 ft) 6. Vorywide (5.20 ft) 7. Entremety wide (2.20 ft)
TA SHEET ge landslide			Method to Determine Compressive Strength 2	DISCONTIN Length of fine (m)	\mathbb{X}	1. Extremely close (<1 ki) 4 2. Very close (1:2.5 ki) 5 3. Close (2.5 - 8 ki) 6 7
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State of Alaska DOT & PF Design & Engineering Services Statewide Materials	Nest Stope			Block	1. Very large (>2161° or 61 cube) 2. Large (>2161° or 2.61 cube) 3. Modemic (>271° or 0.57.21 cube) 4. Smal (1221°-0.31° or 234n 0.051) 5. Very small (<12 n° or 2.34n 0.061)	Jainted
State o Design 8	GENERAL INFORMATION Location 250' West	1. Natural Exposure 2. Construction Eccaration 3. Trial Pk 4. Trianch 5. Adi 6. Tunnel 7. Dial Hole	ROCK MATERIAL INFORMATION Color 2 2 3 4 1.Light 1. Freisch 1. Feit 2. Denk 2. Denk 1. Peit 4. Brownish 3. Yeitow 5. Generation 2. Steven 6. Generation 2. Bio- 7. Blusch 7. Blue 8. Generation 2. Generation 7. Blue 8. Generation 2. Generation 7. Blue	ROCK MASS INFORMATION	1. Biocky 2. Takutar 3. Columnar 4. Shattered	Failure Mode e.g. Topping, Wedges

Yukon River bridge landslide: Preliminary geologic and geotechnical evaluation

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Appendix A. Forms Effective October 1, 2003

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Field Rock Classification and Structural Mapping Guide

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Field Rock Classification And Structural Mapping Guide

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Rock Type Greet unacke	Grain Compressive RG Method to Determine Size Strength 2	ROCK MATERIAL INFORMATION
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tion, design issues etc.) Saming gruck ~151 feet escon upper 1/4 of	REMARKS (exposure type/age, stability condition, design issues etc.) StationHole No. 掛けり No. of Sheets of No. of Sheets of The 250' Station. No. of Sheets of Discontinuity Data 200 Control to Active to Active that makes up upper 1/4	GENERAL INFORMATION Location (250) 1,003/ 05 b-17c Locality stope
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sheet 2 of 2

Yukon River bridge landslide: Preliminary geologic and geotechnical evaluation

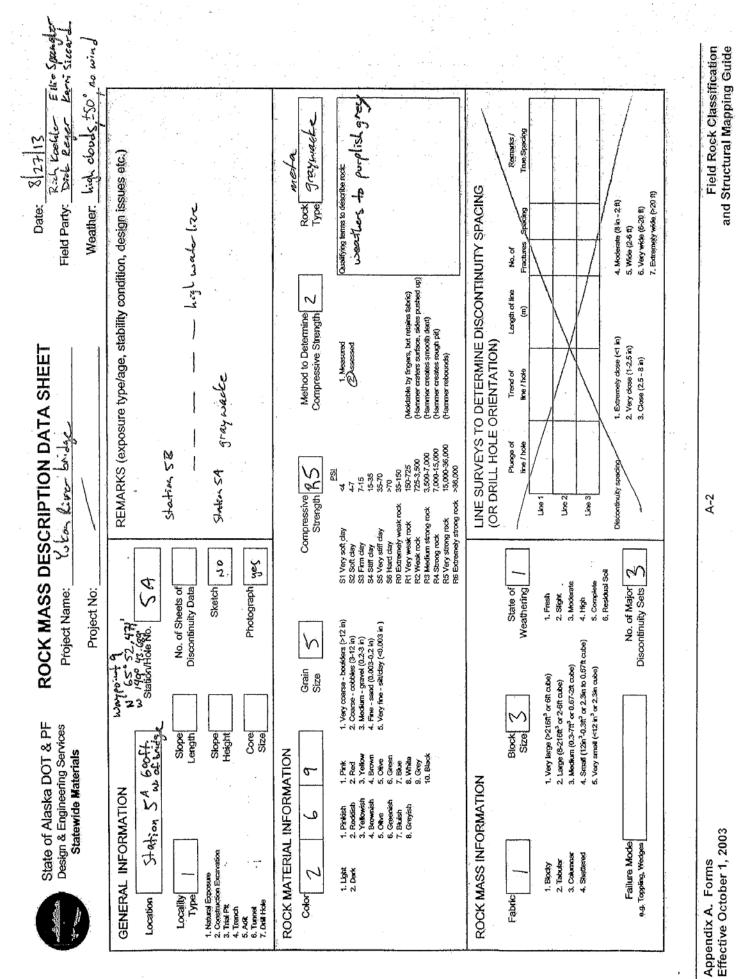
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ЕҮ ИАІА әпес і ид			Surface Waviness Waviness Shape Wavelength Amplitude	1		a state of the second se	6 K	2 12 -54	12)	0 12" 5"	12" .5'				1/11 - 2m	- a		2	COMPRESSIVE STRENGTH OF INFLUING	PSI S1 Very soft day 44 S2 Soft day 4-7 S3 Fam day 7-15	24 Stiff day 15-35 55 Very stiff day 35-70 36 Hard day >70		R2 Week rock 7.25,500 R3 Medium strong rock 7.300-75,000 R4 Sheng rock 7.000-75,000 R5 Very strong rock 15,000-35,000	1	
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	f bridge	ON OF DISCONTINUITY	Aperture/ Persistence Termination Width	-	2					00	2	2	2	2	č			-	a	A TEXT UNCLIMIN 1. UNC 1924 (< 2004 m) 2. Tagat (.00401 m) 1) 3. Party Open (.0102 m) 4. Open (.0211 m) 5. Uncorrective Value (.1.4 m)	<u> </u>	s, Carecrous (> 3 ft)	SHANESS SURFACE SHAPE 1. Planar 2. Undutating	3. Stepped	
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Field Rock Classification And Structural Mapping Guide

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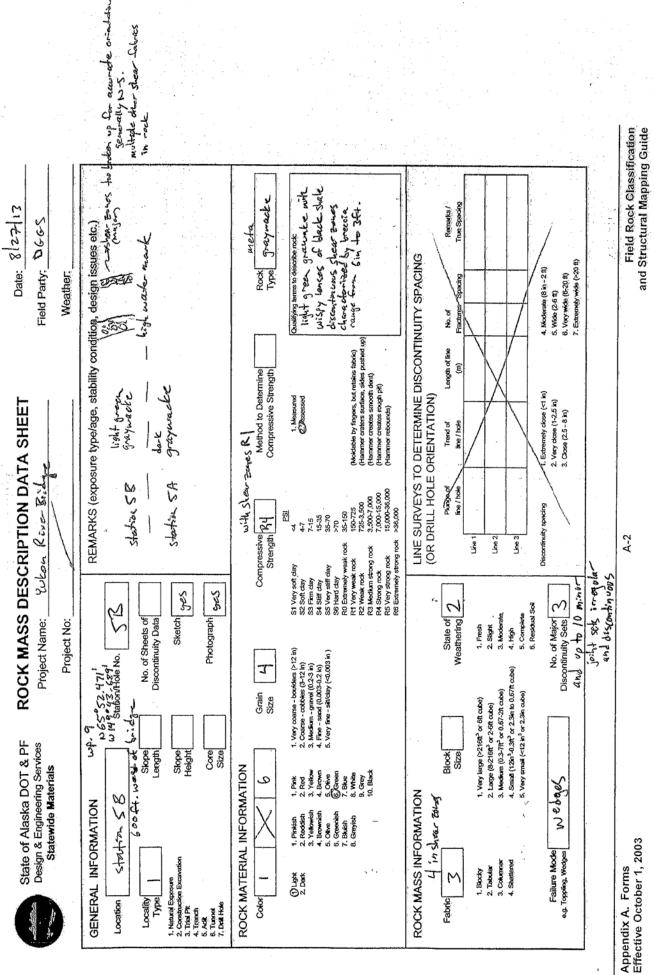
Appendix A. Forms Effective October 1, 2003

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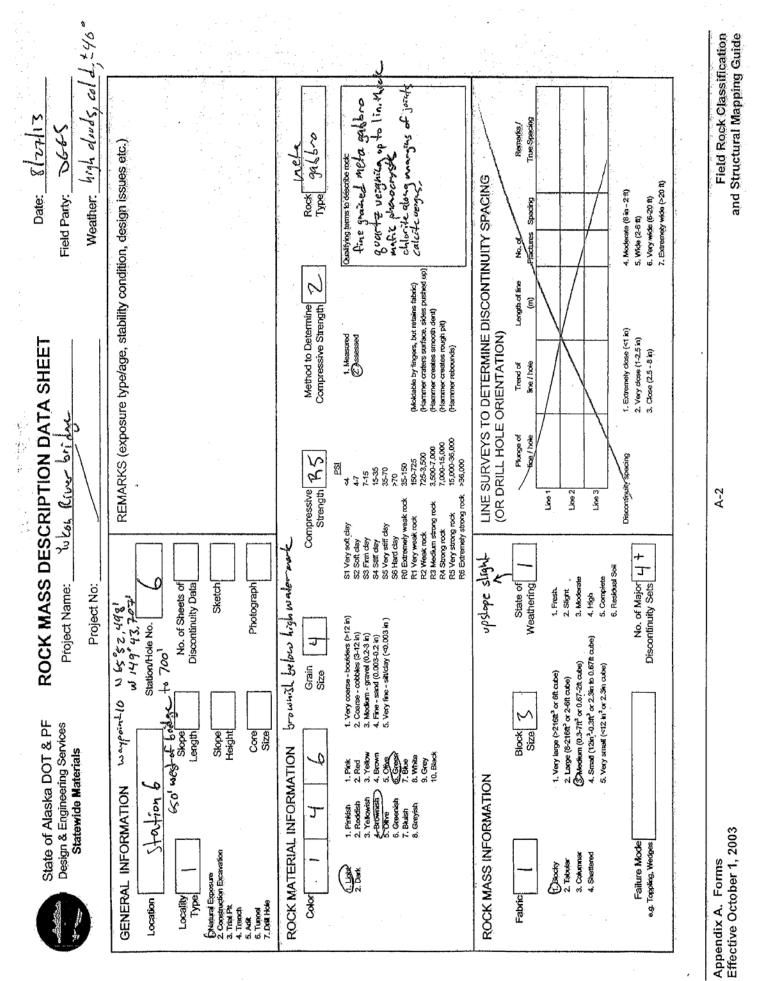


Yukon River bridge landslide: Preliminary geologic and geotechnical evaluation

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U	GENER Location:	ITAN	Station	or Depth					head.	Trobace new 1		>	Aon o D								TYPE 0. Fault	1. Fault 2. Joint	3. Cleanage	5. Shear	6. Fasture 7. Tension	8. Folkation 9. Bedding		1ECMI	1. One 2. Both	-			

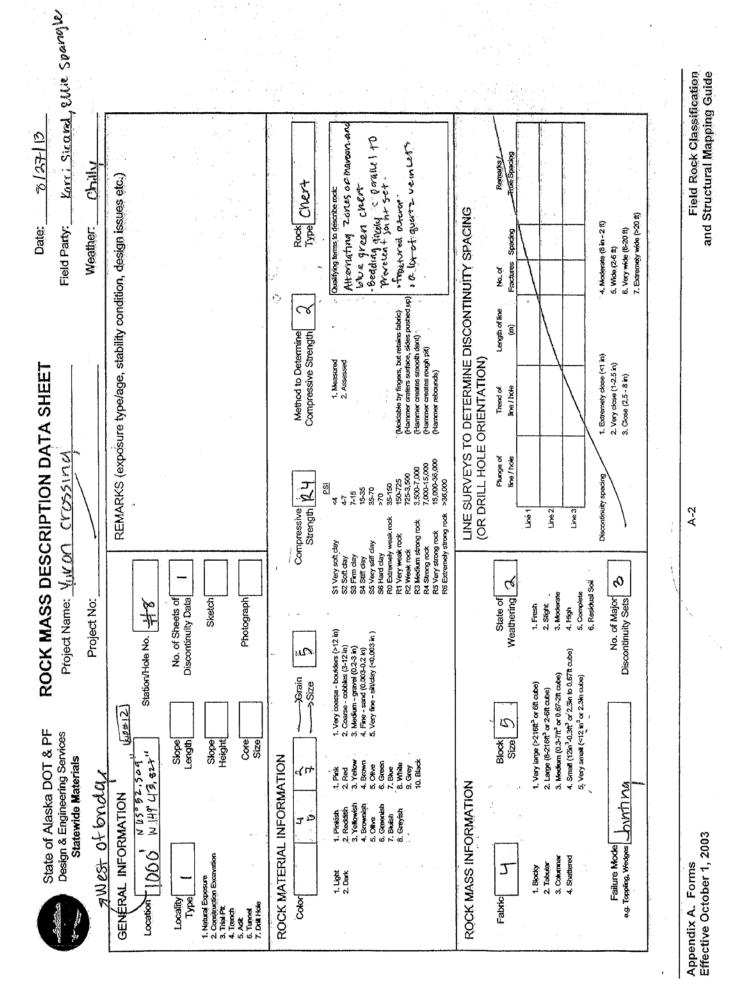
Appendix A. Forms Effective October 1, 2003

A-3

Field Rock Classification And Structural Mapping Guide

		DUCK MASS DES	DESCRIPTION DATA SHEFT	A SHFFT	Date: 8/27/2013	
	Design & Engineering Services		Yuken Crossing		Field Party: Karni Sicarol	ellie Spangler
	Statewide Materials	Project No:				
GEN	GENERAL INFORMATION		REMARKS (exposur	REMARKS (exposure type/age, stability condition, design issues etc.)	tion, design issues etc.)	
Location	tion 230 to 800 Wet	Station/Hole No.	5	examined at this	examined at this station is located	
	Locality Slope Locality Type	No. of Sheets of Discontinuity Data	To salah Dan sa	- Jun and - LO		· · · ·
1. Natural 2. Constru 3. Trial Pt	1. Natural Exposure 2. Construction Excavation 3. Trial Ph. 4. Trial Ph.	Sketch				
5. Add 6. Tunnel 7. Drait Hole	od Core	Photograph				
ROC	RIAL INFORMAT		Compressive Q 氏	Method to Determine	Rock Dror ite to Gablarty	(see this section).
-	4 4	Size 1	Strength 13		Type	
	1. Pink 2. Red 3. Yekow 4. Brown	1. Very coarse - boulders (>12 in) S1 Very soft day 2. Coarse - oxbbas (3-12 in) S2 Soft day 3. Medium - gravel (0.2-3 in) S5 Firm day 4. Final - sand (0.03-0.01) S5 Vint day 5. Vonc 6 salvines (n) S5 Vint day	day 44 7-15 15-35 4ev 35-70	1. Measured 2. Assessed	Qualitying terms to describe rock. Fine grained, phareers - Quarts, fe bispar, bistiftes.	
	Greenish 6. Green Bluish 7. Blue Greyish 8. White 9. Grey	i,	>70 35-150 150-725 725-3,500	المرادلية المراجع ا مراجع المراجع ال	uninhoue . Some ventry Gasterer , guart "of have mental and these	
	10. Black	R3 Medium strong rock R4 Surger rock R5 Eutremety strong rock	3.500-7.000 7.000-15.000 15.000-36.000 * >38.000	(Hammer creates smooth dent) (Hammer creates rough pit) (Hammer rebounds)	-Oxde staining	
ROC	ROCK MASS INFORMATION		LINE SURVEYS TO DETERMINE (OR DRILL HOLE ORIENTATION)	LINE SURVEYS TO DETERMINE DISCONTINUITY SPACING (OR DRILL HOLE ORIENTATION)	NUITY SPACING	
_ ŭ	Fabric 1 Block 3	State of Weathering	Plunge of time / hole	Trend of Length of line Bre / hole (m)	No. of Rematra / Fractures Spacing True Spacing	
	1. Blocky 1. Very large (>216f ¹⁵ or 6ft cube) 2. Tabriar 2. 1 anno (8.216ft ² or 2.6ft cube)	or 6it cube) 1. Fresh An crive) 2. Skort				
	* *	t cube)	Line 2			•
-14	6. Very smell (<12 kt ³ or 2.3kn cube)	or 2.3in cube) 5. Complete 6. Residual Soli	1			
	Faiture Mode	No. of Major Discontinuity Sets	Discontinuity spacing 2 3	1. Extremely dose (<1 ki) 2. Very dose (1-2.5 ki) 3. Close (2.5 - 8 ki)	4. Moderate (3 in – 2 it) 5. Wee (2-6 it) 6. Very wee (-20 it) 7. Externedy wee (-20 it)	
5. 			au - 1		- The second se Second second sec	
Appendix A. Effective Oct	Appendix A. Forms Effective October 1, 2003		A-2		Field Rock Classification and Structural Mapping Guide	ification g Guide

\$kari Sicard				at styles	Telorake, gta Fill w/cllline	Eta fill whereas	STESTICK wich , rake Sidon	RAN - 124/115 - QTZ, BLUE 505	-	Corparte terminarian	DINN 37 NO SOS	19				d	Ward K wall to be cuspate				-					actual, Ru	 	Forms 1, 2003
Date: 8 27/13 Field Party: Ellie Sounder 48 Weather: Sunny ~ 55°F	Disontinuity Data Sheet No.;		Flow AZ Re	1 198 4		++	0 359 4 N 205 4	054 l	040	1 5 (34215 HACK	OHS 3	1-2-1		1.	0 248	6 207 S	SVEC 1 4 200 0		ERFLOW (Open)	upscontanty vor jugatom dy, stater som over a prosense jescontinuely dy, to eyklence of water flow Discontionely dy, storys endonce of water flow e.g. rust stariging	Discontinutly is damp, but no the water present Discontinutly shows seepaga, occasional drops of water, no continuous	continuous flow of water (Estimate Umin and describe pressure i.e. kow dum high)		ting mate	mater Mmit	Next changes (Estimistic limit and describe pressure)		Appendix A. Forms Effective October 1. 2003
JRVEY DATA SHEET			Surface Wavness Waviness Shape Wavelength Amplitude			3	7-7 					1					7 10 5		COMPRESSIVE STRENGTH OF INFILLING WA	2 v	47 7-15	Set Stiff clary 15-35 thome 25 C SS Very stiff clary 35-70 5, G St Part clary 75 25-70 6, G	35-150 150-725	R2 Weak rook 1/25-5,500 006 R3 Medum strong rock 3,500-7,000 7. F R4 Strong rock 7,000 8. F	15,000-35,000	1 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1		A-3
Project No: און אין אין אין אין אין אין אין אין אין אי	Station/Hole No.:		Aperture/ Nature of Strength of Surface Width Failing Failing Roughness	11	8942	4 8 4 × 1 × 1	8 at 2	· .	18-42	1 50-7-2	2202 0	4 3 Sect 7 79	1 2 3	-]	 	0	8012 KJ	NATUREOF FILLING	1. Clean 2. Surface staining 3. Mon-colorestyre	4. Inactive 5. Swelling	5. Cemente 7. Critorite. 8. Other - 5	-	50	5 10 Smooth, un 15	20 Rough, stepped		A
State of Alaska DOT & PF Design & Engineering Services Statewide Materials	GENERAL INFORMATION LOADON \$400 West	DF DISCONTINU	Station Dip Dip Persistence Termination	M OLC IL C		5 73 90 E 3 0	85 90 E 3	ce h	30 25 56	100 5 3	39 186 5 3	-	2 TX ON 2	2 34 0 2	2 84 =0 MM 2 2	2 37 00 2	6N 2	36 0 N 1	PERSISTENCE	# Zone 1. Very Low (~3 ft) 2 Low (3-10 ft)		Shear Fissure Torreion Crack	S. Carrienton S. Carrientous (> 3 ft) bodding		1. One end vettile 2. Smooth 2. Undutating 2. Both ands visible 3. Rough 3. Shepped	3	N 146- 40- 1-1-1	Field Rock Classification
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Yukon River bridge landslide: Preliminary geologic and geotechnical evaluation

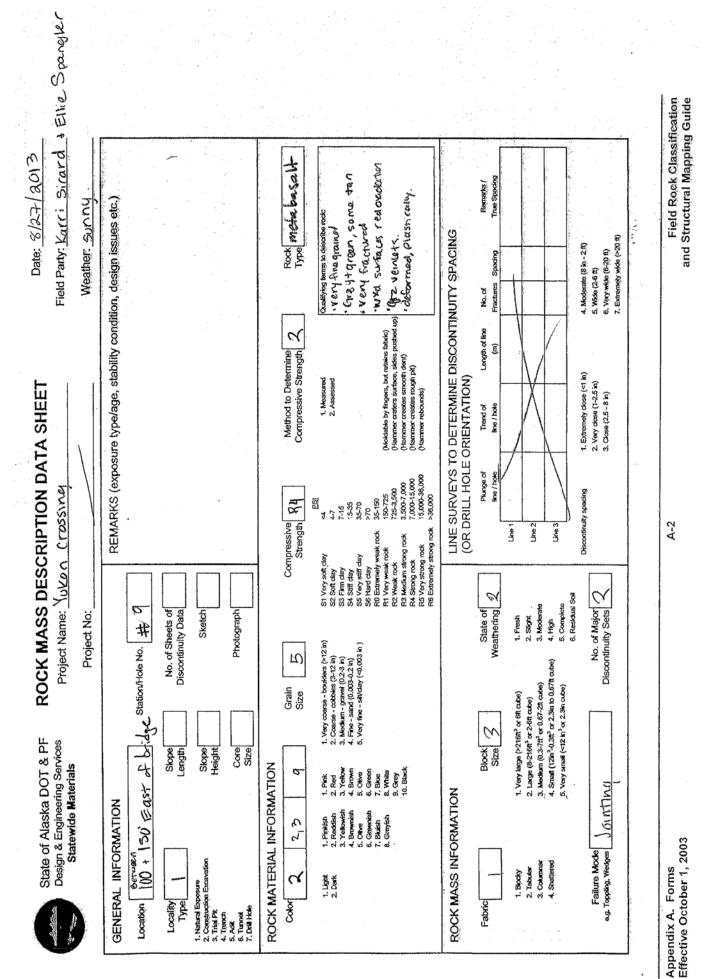
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Appendix A. Forms Effective October 1, 2003

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Field Rock Classification And Structural Mapping Guide

Preliminary Interpretive Report 2013-6



Elen Spar			K	225		· •		P .						-	- - -	1	30			0,		IN U -			-				
Field Party: <u>Karn' Sirond & C</u> Weather: <u>Sunny</u>	Disontinuity Data Sheet No.: 7 of 7		JRC From A2 1 Remarks Sectors	0 030	0	0 236		<u>0</u>	0 056	0 055 1	0 254	0 230	0	0 1001		0 121 6 2QU	0	0		W WAY LOW VAM N.	8 0 235 WATED BLOW (Press)	WALLEX FLOUR (Appendix) 0. Discontinuity (way tight and day, water flow docsn't appear possible 1. Discontinuity day, no evidence of water flow 2. Discontinuity day, shows entitience of water flow e.g. rust stanting	 Discontinuity is damp, but no thee water present Discontinuity shows seepage, occasional drops of water, no continuous 	sow. 5. Continuous forw of water (Estimate limin and describe pressure i.e. low modum high)	WATER FLOW (Filed)	or many instruction province under present of the second	 Filling meterial wet, occasional drops of water Filling meterials show stors of outwash, continuous from of water 	(Estimate intraction in the intervention out, considerable water flow along rot. Fifting materials locally weshed out, considerable water flow along outwark intervets (Estimate light) and describe pressure)	
			Strape Wavelength Amplitude																			ESSIVE STRENGTH OF INFILLING PSI soft day 44		tay 15-35 stiff clay 35-70 clay >70			5.5	R6 Edmenely strong rock >36,000	
Yukan Crossind	# 9 Harrison		Strength of Surface Surface Filling Roughness Shape		- 2		40	1 6	ь 1	4		*	1 8×10	×1 8	A 4	1) () e	6	R5 1	d	S CING	r clay matrix x clay matrix	6. Cemented 5. S4 Stift day 7. Chloride, talc or gypsum 55 Very stiff day 8. Other - seedty 58 Herd day					
Project Name: <u>Yuv.cn</u> Project No:	Station/Hole No.	LINNI	Aperture/ Nature of Width Filling	-		-						-	8	-	-				4	4 Giorz	ł							10 Smooth, unoutaing 15 20 Rough, stepped	
ring Services aterials	warpoint 13	ON OF DISCONT	Persistence Termination	2 1	- 2		2		1 4	-	- 4			4		2		× (1.	67	₩.	APERTUR 1. Very 5gk 2. Tight (J 3. Partiro			9. Caremous (> 3	CHNESS SURFACE SHAPE		3. Stepped	
Design & Engineering Services Statewide Materials	GENERAL INFORMATION COMPAND	NATURE AND ORIENTATION OF DISCONTINUITY	Dip. Time Din Direction	510	ž	30	- 1051	-	9 L.A. A.C.	F	14	4	S I	34	50	20	53	(N)() CL 7	10	200	57	PERSISTENCE 1. Very Low (53 ft) 2. Low (5-10 ft) 3. Lowform (10.20 ft)	4, High (30-50 ft) 5 Very High (26-01)	:	ť.	SI IDEXCE BOUICHNESS			
Ð	GENERA	NATURE		or Lepun			3			<i>x</i>	-			5	2	2			4 4	6 5		TYPE 0. Fault Zone 1. Fault	3. Cleanage	5. Shear 6. Fissure	/. rension used 8. Foliation 9. Bedding		0. Neither and visible 1. One and visible	2. Both ends via	

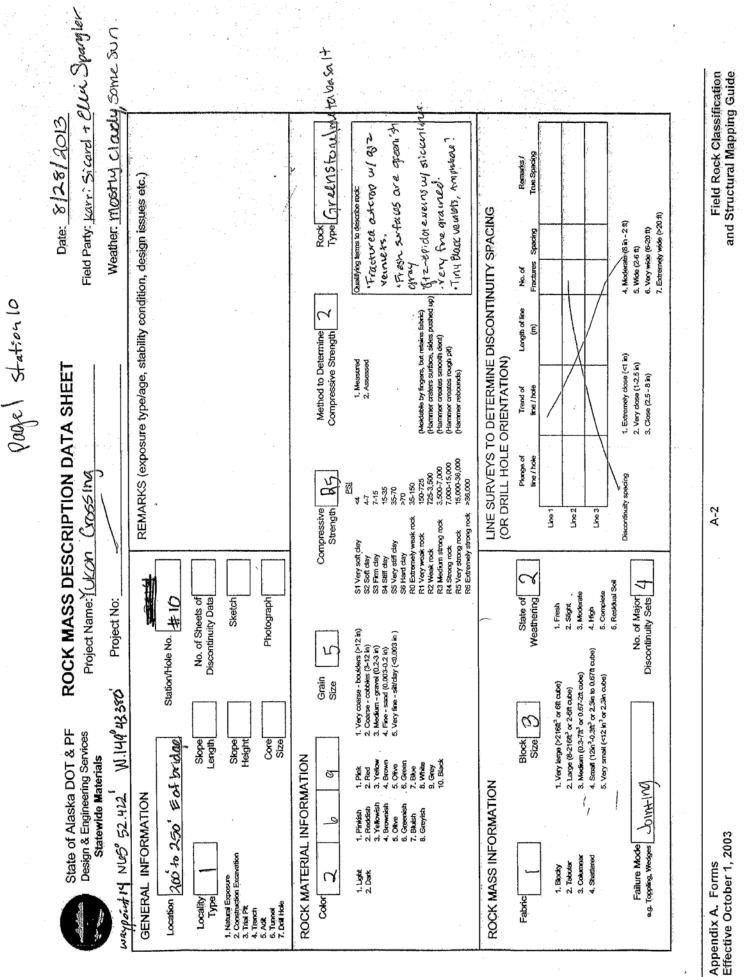
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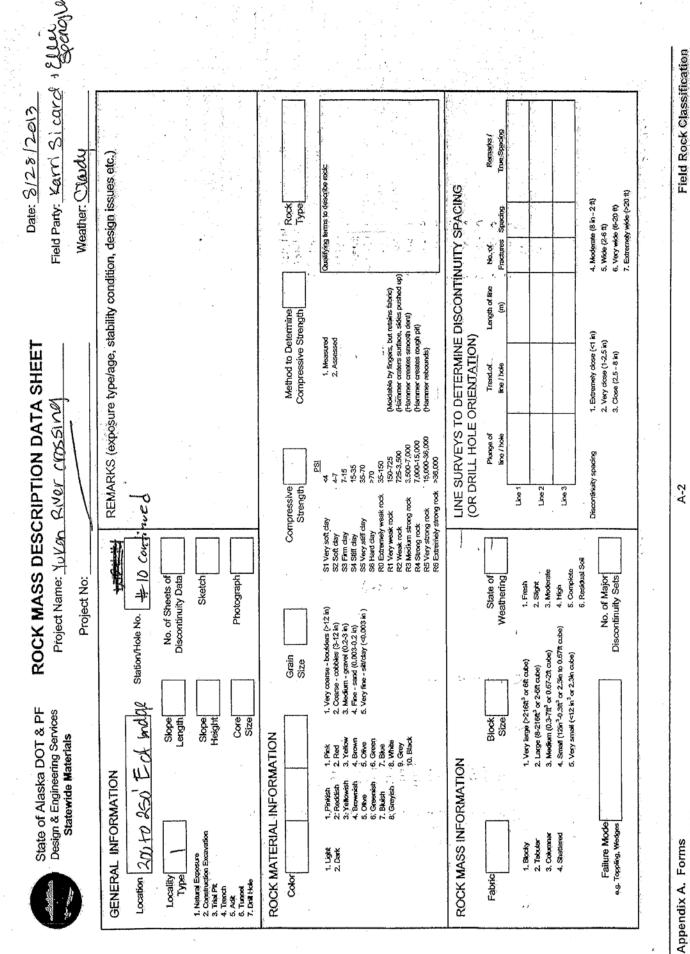
Appendix A. Forms Effective October 1, 2003

> Field Rock Classification And Structural Mapping Guide

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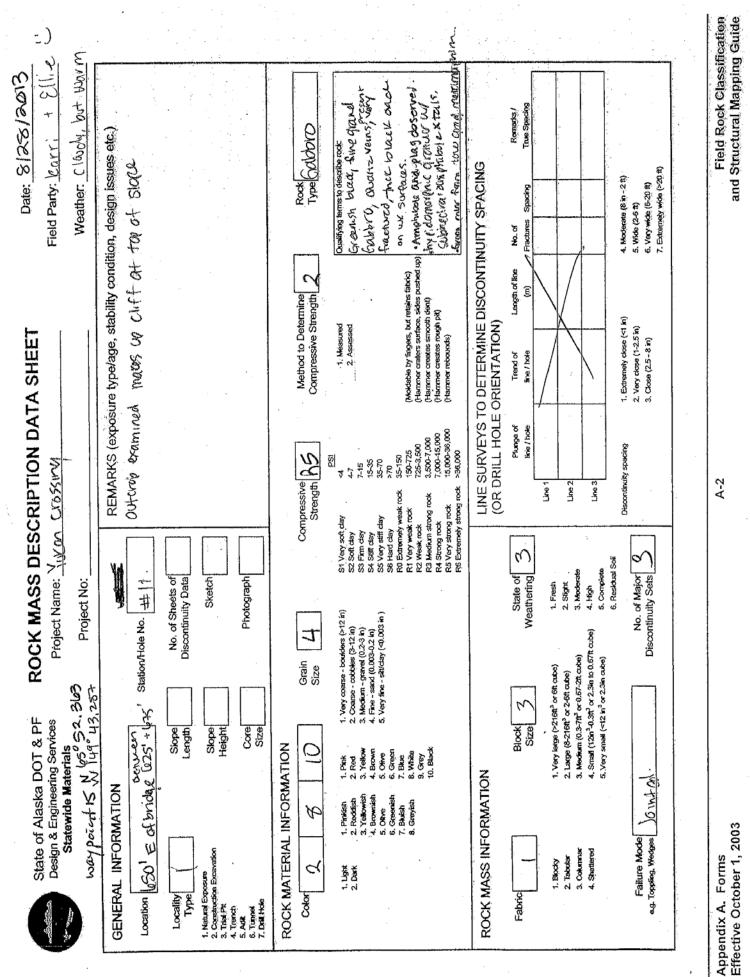
Page 2 Station 10



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and Structural Mapping Guide

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Field Party: KANT		Disonanury Data Sheet No.		JRC Flow AZ	3 INT IFIENCE	000	0000	22	0	296 5 of	0 243 5	C 302 5 of tu	224 V	0 212	0 204	322 C	0 343 942 0	> 344 246	- 0 2184 30% PM	2 27 11 20	13 0 04 11 10 23	WATER FLOW (Open)	 USCONTRINY YAY BUT AROUNTLY, WARA AND Discontinuity dry, no evidence of water if Discontinuity dry shraws evidence of wat 	is dam	Row. 5. Continuous flow of water (Estimate findin and describe pressure i.e. tow 			Νœ́	 Funds materials show sights or outwass, usuations new or neural (Estimate Timin) The material instrains locally wished out, considerable weiter flow along connected character Efformation in description forestres) 			
655 IN 19		60		ress right	2 20" 1.5"					-			40		2 2" 2"	1 2"				11 = C	2 10 11	COMPRESSIVE STRENGTH OF INFALING	PS!		S4 SVF day 15-35 S5 Very stiff clay 35-70			R3 Medium strong rock 3,500-7,000 R4 Strong rock 7,000-15,000	R5 Very strong rock 15,000-36,00 R6 Extrang rock >36,000			
Project Name: Witen River Crossing Project No:		Station/Hole No. 410 Confirme		Aperture/ Nature of Strength of Surface Width Filling Filling Roughness	2.8 22	1	{ }	2 _ 2		2 - 2	2 - 2		x	8/1939 KS -	0	8 air 25 3,1	æ	RS	-	+		OF FILLING		 Non-concerve Inactive clay or clay matrix S2 Summon clay or clay matrix S3 			22.8	JRC (Joint Roughness) R3 0 Sickersifood, planar R4		noddais 'ußinnu' nz		
t		2 Sul Statio	OF DISCONTINUITY	Aperture Persistence Termination Width		- 2				2 ~	2			n+ n 				-		2	+ / + 	APERTUREMIDTH	1. Very tight (<004 in) 2. Tight (.00401 in)			8. Extremely wide (4 in - 3 ft) 9. Cavemous (> 3 ft)		\. [*]	2. Undutating 3. Stepped			
Design & Engineering Services Statewide Materials	GENERAL INFORMATION	Location: Between 200 . 2 2	NATURE AND ORIENTATION OF DISCONTINUITY	Station Dip Dip or Denth Tune Dip Direction Pe	1 72 166	2 63	1180	2 109 D (N)	32	8	18	29 (*		24 0 (N)		1.1	56	54	90 Q.1	5 SH O(N)	35		0. Fault Zonia 1. Very Low (<3 ft) 1. Fault 2. Low (3-10 ft)	2. Joint 3. Modum (10-30 ft) 3. Clearinge 4. High (30-60 ft)		Tension Crack Foliation	Beddling		1. One end visible 2. Smooth 2. Both ends visible 3. Rough			



				Space NV	n	5	0,0	11	01	d'W	M	n	4	20	1	617	8	3	5	2	1	0													
ELLIE Spangle					05				S																		-				-	-			
Field Party: Karr: Scord + Ellie Weather: Choudy but Warrh.15h				A.2. Remarks	2	Strichlorite	334 100 39 art chlorit	Place Original Succes	255 36 25° AGEN ADM NN NO D	-	2600 Rake 4500 from M.	0.38	3	363		AUL Dame Mat Set -	253	93	233	21	1 24	215 1)	(Open) very light and dry, water flow doesn't appear possible	y, to evidence of water flow	 Discontinuity only, shows envisions or water how e.g. rust summing Discontinuity is damp, but no free water present 	 Discontinuity shows seepage, occasional drops of water, no continuous flow. 	Continuous flow of water (Estimate Umin and describe pressure i.e. low exium Nichi)		hoavity consolidated and dry, significant flow unlikely	is a very tom pour commy Filling matterials damp, no free water present	 Filting material wet, occasional drops of water Filting materials show signs of outwash, continuous flow of viater 	(Estimate limite) 10. Failing materials locally washed out, considerable water flow along	outwash channels (Estimate Vinin and describe pressure)		
Eiel		nuosin		Water JRC Flow	0 6		903	~	30	20	0	φ	0	n v 0 ()c	1	х 0 1	0					WATER FLOW	1. Discontinuity dry, no	2. Discontinuity is o 3. Discontinuity is o	 Discontinuity she flow. 	5. Continuous flow medium high)	WATER FLOW (Filed)	6. Faling materials hear				outwash channels		
Project No: Yukon CrosStra				Waviness Waviness Wavelength Amplitude	4																		COMPRESSIVE STRENGTH OF INFILLING	ŝ	4 4	7-15	35-70		150-725		7,000-15,000 15,000-36,000				
CrosSing				Surface Shape	r	3	6		n I	(n	S	m		1				-			-		COMPRESSIVE ST		S1 Very soft day S2 Soft day	SS FIFT CAY	SS Very stiff clay SB Hard clay	BD Extramely weak	Rt Very weak rock	R3 Medium strong n	R4 Strong rock R5 Very strong rock	R6 Edmenely strong rod			
Yukan cras		朱儿		Strength of Surface Filling Roughness		R5 1	1 50	1	R) L	1-1		1	5		Å (10	7	1	5	4	4	2	NATURE OF FILLING	steining	Non-cohesive Inactive clay or clay matrix	5. Swelling clay or clay matrix 6. Cemented	7. Chilomie, taic or gypsum 8. Other - streictiv			ughness)	ed, planer	ndollating	pede		
Project Name: Project Name:		on/Hole No.:	7	Aperture/ Nature of worth Failing					5 2,8 at	44	€	2	2	21	Ye		-	3		-			NATURE C	2. Surface staining	 Non-cohesive Inactive clay o 			R. A E	- 	JRC (Joint Roughness)	0 - Skokenskided, planar 5	10 Smooth, undulating 15	20 Rough stepped		
		20	NATURE AND ORIENTATION OF DISCONTINUITY	Tarmination				0	д,	10	4	Ч	0	S).								APERTURE/WIDTH	2 Tight (.004-01 in)	3. Party Open (.0102 in) 4. Open (.021 in)	5. Moderately Wide (.14 in) 6. Wide (> 4in)	7. Vecy wide (4 - 4 in)	S. Carremous (> 3 ft)		SURFACE SHAPE	1. Planar 2. Undulating	3. Stepped			1. A
e of Alaska DU I & jn & Engineering Servi Statewide Materials	NO	, hetwan (25'+475	VTION OF D	Dip Dimeters			1	3				9	÷.			1 r			4 (m)	2 (1)	(m) 2	2 1					5 / «) Ö		SURFACE ROUGHNESS	3	 			
State of Alaska DO1 & PF Design & Engineering Services Statewide Materials	NFORMATH	O' , lathuan	D ORIENT/		<u></u>	+				AT 104 121	03601	56 1	70 160(5)	Ø	5	02 NN	1.1	+			いのうき	(m)022 2h	PERSISTENCE	2 Low (3-10-11)	3. Medium (10-30 ft) 4. High (30-60 ft)	5. Very High						3. Rough			
	GENERAL INFORMATION	Location: [250	NATURE AN		or Leptin Type	1	10	3	" "	5	220	5	10 2	50	2	500	-	-	2	5 1 2	رد ا ک	2 2	TYPE	U. Faunt Lorie	2. Joint 3. Clearade	4. Schistosity	6. Fissure	8. Foliation	9. Bodding	TERMINATION	0. Neither and visible 1. One and visible	2 Both ends visible		•••	

	3	idy													fication 3 Guide
Date: 8/28/2013	Field Party: Kow & FULU	Weather: Sunny t cloudy	ttion, design issues etc.)						Rock Type Diorite	Qualitying herms to descarbe made - Ogre, felds gov, Am philipole - Fine graunel - fire vel ns - fractured (hean kr) - hypictomorphic gram un sub-edroct any philipole	NUITY SPACING	No. of Remarks/ Fractures Specing True.Specing		4. Moderate (8 ři - 2 ři) 5. Wicke (2-8 ři) 6. Very wicke (5-20 ři) 7. Estresnešy wicke (5-20 ři)	Field Rock Classification and Structural Mapping Guide
DATA SHEET	0		REMARKS (exposure type/age, stability condition, design issues etc.)		ł.				Method to Determine Compressive Strength	1. Messured 2. Assessed (Modatile try fingers, but retains fabric) (Hammer craeties surface, sides pushed up) (Hammer creates smooth den) 00 (Hammer retexends) 000 (Hammer retounds)	LINE SURVEYS TO DETERMINE DISCONTINUITY SPACING (OR DRILL HOLE ÒRIENTATION)	Plunge of Rend of Length of line line / hole line / kole (m)		1. Extremety dose (<1 in) 2. Very dose (1-2.5 in) 3. Close (2.5 - 8 in)	
ROCK MASS DESCRIPTION DATA SHEET	LUKON CYDSSIN	7	-						Compressive R5 Strength	PSI S1 Very soft clay 4 S2 Soft clay 4 S2 Soft clay 47 S2 Soft clay 47 S3 Fam day 7-15 S4 Stift clay 7-15 S4 Stift clay 7-15 S4 Stift clay 7-15 S5 Very stift clay 35-70 S5 Hand clay 7-16 R0 Extramely weak rock 35-70 R1 Very weak rock 35-70 R2 Weak rock 755-3,500 R3 Medium strong rock 15,000-36,000 R4 Extramely strong rock 15,000-36,000 R5 Extramely strong rock 15,000-36,000	LINE SURVE	Ptun	Line 2 Line 2	Sol Discontinuity specing	A-2
ROCK MASS	Project Name:	Project No:	223	vHole No.	No. of Sheets of Discontinuity Data	Sketch	Photograph		Grain U Size	1. Very coarse - bounders (>12 in) 2. Coarse - cotches (3-12 in) 3. Medicim - gravet (0.2-3 in) 4. Fine - sand (0.003-0.2 in) 5. Very fine - salv(day (<0.003 in)) 5. Very fine - salv(day (<0.003 in))		State of Weathering		6. Resoluted Source No. of Major	
State of Alaska DOT & PF	Design & Engineering Services	waybourt 16: N lo 5°52.338	GENERAL INFORMATION W 149.43.223	E - Berwan	et to the Stope	Stope	Core	VFORMATION		1. Prinkish 1. Park 1. Very 2. Reoddish 2. Red 3. Yebows 3. Mook 4. Rrownish 4. Brown 4. Fine- 5. Other 5. Other 5. Very 5. Other 5. Other 5. Very 7. Bush 7. Blue 8. Geryish 9. Gery 10. Blact	RMATION	Block	1. Very large (>2161° or 61 cubo) 2. Large (8-2161° or 2-61 cubo) 3. Medium (0.3-711° or 0.67-21 cubo) 4. Small (122n ³ -0.31° or 2.31n to 0.671 cubo) 5. Very small (<12 h ³ or 2.31n cubo)	Jointed.	
State of /	Design & I	triogram	GENERAL INFORM	Location NICK \$50 E	Locality	1. Natural Exposure 2. Construction Excavation 3. Trial Pt. 4. Trench	7. Dout Hole	ROCK MATERIAL INFORMATION	Color		ROCK MASS INFORMATION	Fabric	1. Blocky 2. Tsbular 3. Columnar 4. Shaffared	Faiture Mode	oendix A. Forms ective October 1, 2003

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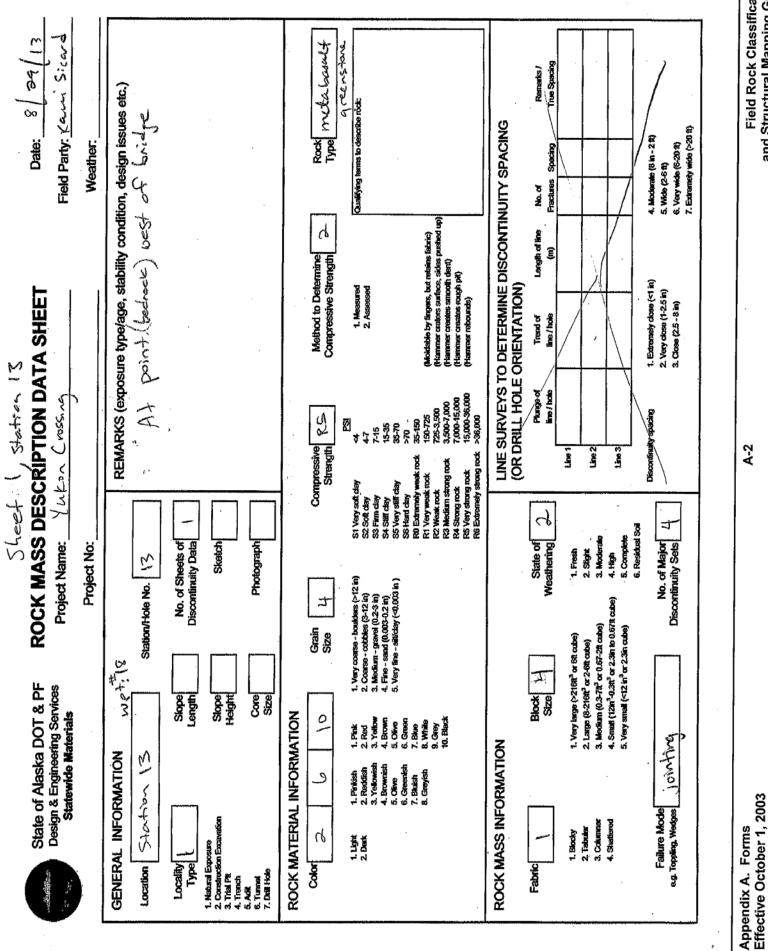
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Field Party: <u>Larn: S: raid</u> . Weather: (<u>\oudul</u>	. [Disontinuity Data Sheet No.:		ats Spacifiel	A Steres 1											liake of 45° downsom			1 8 x' 9 + 2 1	8	Rake Se down toma	10MC	0. Discontinuity very tight and dry, water flow doesn't appear possible 1. Discontinuity dry, no evidence of water flow	Discontinuity dry, shows evidence of water flow e.g. rust stanting	 Discontinuity is camp, but no new weak preserve. Discontinuity shows seepage, occasional drops of water, no continuous 	flow. 5. Continuous flow of water (Estimate trimin and describe pressure i.e. low		WATER FLOW (Filed) 6. Failing materials heavily consoliciated and dry: significant flow unlikely	due to very kow permeability 7. Failing materials damp, no free water present	 Filting melerinal wet, occassional drops of water principal melerinal wet, occassional drops of water principal meterinal show signs of outwash, continuous flow of water features limits? 	countrie within) 10. Falling materials locally washed out, considerable weter flow along outwash channels (Estimate limin and describe pressure)	· ·			A nhandiv A
eld Party: Weather:		nuity Data		AZ Remarks	, i	212	234	1 Oalt	10	و	237	2 ~~~	1	135	2510	-	5/07	324		086 154	Con lico	5	 Discontinuity very tight and dry, water flow in Discontinuity dry, no evidence of water flow 	, shows evide	we seepage,	of water (Est		ied) heavity conso	meability damp, no free	et, occasion show signs o	s localty washer (Estimate Vimin			4	
Fiek V		Disontir	· ·.	Water A	0 200	+	00		6			2	21		0	-	-		-			Ś	ontinuity ver ontinuity day	ontinuity dry	ondioutity sho	mous flow	(hội lungh)	WATER FLOW (Filled) 6. Filling materials hear	due to very low permeability 7. Faing materials damp, n	g material w g materials de l/min)	(countries units) 10. Faling materials loca outwash channets (Estin				
· ·				JRC	Z	=		=	N	m	~ ~	v	1	5	'n	<u> </u>	a	m	7(- 1-	0 1	WATE	0. Disc	2 Diso	3. Diso	5. Corr	medium	WATE 6. Fillin	due to	1991-100 1991-100 1991-100 1991-100	10. Fill				
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SHEET					engin Amplitude			-		12.		11 11 11		17.11	10 11 11	(((((()	GTH OF INEN 1 INC	8	24	7-15 7-15	87-18 19-18 19-18		725-3500 3,500-7,000		
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Appendix A. Forms Effective October 1, 2003

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Project Name: Project No: Project No: Project No: Project No: Project No: Project No:	はたった たたった Startogit of Surfaces 用語の Reling Reli	B Surface Warehength	Mawiness Amplitude		Field I Field I Water Water Notation Notation	Field Party: Carry Sicard Weather: overcess + ~55 ° F Weather: overcess + ~55 ° F Mainity Data Sheet No:: 2.0 0 0/55 1 32.0 2.15 2.15 2.15 2.15 2.0 2.15 2.0 2.15
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4. Open (.021 in)	r clav metrix	aly sourcesy offician	1 1	Z. USCO	conucy dry, s	2. Lissonnantary or y, shows evidence of water flow e.g. rust staining 3. Discretivity is down. but as for units
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0. 1140 (2.441) 7. Vervinida (4.445)		# clay	15-35	HOW.		
8. Extremely wide (4 in - 3 ft) 8. Other - specify 9. Caremous (> 3 ft)		S6 Hand clay	Q.~	o. commuous medium high)	high)	 Contractions frow of water (Estimate limit) and describe pressure i.e. low medium high)
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15 20 Rouch strenged				10. Falling	g materials to	(Longing materials locally weshed out, considerable water flow along

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